

#### http://doi.org/10.54499/ERA-MIN/0002/2019 https://mostmeg.rd.ciencias.ulisboa.pt/



The W-Sn quartz lodes of Mata da Rainha and related anomalous tourmalization; insights from integrated fluid inclusion analysis and boron isotopes

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# Mata da Rainha





Romão et al., 2010

Granite type S : 310 - 295 Ma (Villaseca et al., 2011; Roda-Robles et al., 2018)

- W deposit exploited by Germans during the second World War
- extended zones of tourmalinisation at the contact between schists and the Orca granite
- tourmaline also associated to greisens
- tourmaline in W quartz veins from the deposit
- $\rightarrow$  objectives: tourmaline crystal-(geo)-chemistry and comparisons
  - $\rightarrow$  Nature of the fluids





#### Sample location in the Mata da Rainha area

# Materials and Methods





# Tourmalinites, quartz-tourmaline veins, greisen with tourmaline





Utz

#### Expression of Tourmaline at Mata da Rainha

- schist tourmalinisation around syn-deformational quartz veins
- tourmaline in mineralized quartz veins from the deposit
- tourmalinized schists cemented by quartz
- tourmaline in and at the contact with greisens

# Tourmalinites, quartz-tourmaline veins, greisen with tourmaline



Euhedral tourmalines in thick quartz veins





Tourmalines in needles in quartz veins.







Tourmaline as first vein infilling Growing perpendicular to fracture edge And then in between quartz generations Poorly zoned, probably synchronous of muscovite selvage

#### Tourmalines in micaschists



# Tourmaline zoning - Mapping with micro-XRF



Variations mostly for Fe, Ti (Mg, Mn and sometime Ca).

Ti, Fe and Ca are more concentrated in the outer dark bands.

Mg is more concentrated in the core of zoned tourmalines.











#### Quartz-tourmaline vein







Schist tourmalinization-fine grained zoned tourmalines, « Classic » Fe-Mg zoning Later microveinlets with quartz and coarser tourmalines















**Pervasive tourmalinization** 

MR9Z1B-P2

#### MR 21 A

MR21 A

Quartz –tourmaline vein











- Fe

MG



## Trace elements in tourmalines

The dark growth bands are enriched in MgO (and  $Na_2O$ ).

The light bands are enriched in FeO and TiO







Zn and Mn remain constant at 800 and 1000 ppm along the profiles.

Li shows the same trend as Ti.

Other elements show concentrations of 100 ppm and less.

# Chemical composition of Tourmalines



X (Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>,  $\Box$  = vacancy),

Y (Fe<sup>2+</sup>, Mg<sup>2+</sup>, Al<sup>3+</sup>, Li<sup>+</sup>)

et le site W (OH<sup>-</sup>, O<sup>2-</sup>, F<sup>-</sup>, Cl<sup>-</sup>) (Henry et al., 2011).



#### Tourmalines Mata da Rainha - Argemela - Panasqueira





## Trace elements in tourmalines



**Comparison Mata Da Rainha and Panasqueira** 

Overlay

schists

quantities

in both zones.

on

of

No simple relationships could be defined between the various trace elements analysed. Values similar to those for Panasqueira tourmalines.

1000,0 1000,0 **MR 10** MR 9 between the MR 8 **MR 21** data in the diagrams indicates that the same processes are involved 100,0 100,0 Sn ppm Sn ppm Carocci et al. 2020 Carocci et al. 2020 Influence of the host 10,0 10,0 the trace elements incorporated Codeço et al. 2020 Codeço et al. 2020 into tourmaline. 1.0 1,00 10,00 100,00 1000,( 0,000 0,001 0,100 1,000 0,010 10,000 1 000, 100,000 Li ppm W ppm

# Origin of trace elements in tourmalines



• The composition of tourmalines is partly controlled by the nature of the altered rocks and the physico-chemical conditions of formation.

• Systematic zonation of major element concentrations (Mg, Fe, Ti, Ca)

The succession of dark and light bands on a micrometric scale, linked to the availability of elements in the percolating fluids. (Norton and Dutrow, 2001 ; Dutrow and Henry, 2011).

• Enrichment of fluids with Fe and Ti as in the case of magmatic tourmalines (Marks et al., 2013; Da Costa et al., 2014; Fischer et al., 2023).

The low Sb/Pb ratio would also suggest a magmatic origin.



Beckett-Brown et al., 2023

# Boron isotopic composition of tourmalines

•MR21 (quartz-tourmaline and greisen) and MR10 (quartz vein)  $\delta^{11}\text{B}$  between -10 and -13 ‰).

• MR8 (schists)  $\delta^{11}$ B : -7 to -13 ‰ similar to the tourmaline-bearing rocks of Argemela, but larger range

•Magmatic Fe-rich tourmaline from Segura dykes are the only ones showing distinct values

Echantillon	Faciès	Nombre mesures	δ <sup>11</sup> B (‰) Minimum	δ <sup>11</sup> B(‰) Maximum	Moyenne	Mode
MR21I	granite	58	-13,56	-10,35	-12,10	-11, 34
MR10	vein	8	-12,91	-11,69	-12,50	-
MR8	Host rock	59	-14,00	-6,76	-11,22	-8,84
ARGEMELA 32	Host rock	7	-11,23	-8,68	-10,41	-
ARGEMELA 1	Host rock	9	-12,89	-10,40	-11,18	-
SEGURA 2-1	Host rock	8	-10,01	-8,81	-9,50	-

Segura

Argemela

Tur (granite) Mata da Rainha Tur (quartz veins) Mata da Rainha Tourmalinite Mata da Rainha



# Boron isotopic composition of tourmalines

The tourmalines (quartz-tourmaline and greisen veins) were developed on a hydrothermal edge of the magmatic-hydrothermal system associated with the Orca granite.

The very strong tourmalinisation of the schists produces tourmalines whose isotopic characteristics are slightly more extended but centered on the same values.

At Panasqueira according to Codeço et al., (2017), a significant amount of boron is required. For Codeço, this implies a magmatic source. But remains to be demonstrated.

Values of  $\delta^{11}B$  around -10/ -11 ‰ are effectively close to those of magmatic tourmalines and S-type granites.



Ribeiro da Costa et al 2013 ; Codeço et al., 2017



# Fluids in quartz lodes







Aqueous carbonic fluids

Lc-w : abundant, size <  $30\mu$ m, 3-phases (LH<sub>2</sub>O, LCO<sub>2</sub> et VCO<sub>2</sub>, vapor phase 20 to 40 %).

Lc and Lc-(w): few, size <  $30\mu$ m, vapor phase (90 to 95%) ± water coating

Lw-c: abundant, size <  $30\mu$ m, lfs 2 phases (LH<sub>2</sub>OVCO<sub>2</sub>, vapor phase de 10 to 40 %).

Aqueous fluids

Lw less abundant size :10-30µm. 2 phases, vapor phase 15 to 40 %.

## Fluids in quartz and tourmaline









#### Aqueous carbonic fluids

Lc-(w) and Lc-w : very few, size <  $30\mu$ m, 3-phases (LH<sub>2</sub>O, LCO<sub>2</sub> et VCO<sub>2</sub>) vapor phase 20 to 45%.

Lw-c : **abundant**, size <  $30\mu$ m. 2 phases (LH<sub>2</sub>O VCO<sub>2</sub>) vapor phase 15 to 45%.

#### Aqueous fluids

Lw : **abundant**, size 10-30µm. 2 phases, vapor phase 15 to 30%.





**Quartz lodes** 

Lc-w

Lw-c Lc-(w)

Tourmaline

 $\diamond$ 

#### **Aqueous-carbonic fluids**



## Fluids in quartz and tourmaline







#### Salinity and minimal trapping temperature of the fluids



#### **Evolution in three stages**

1) Crystallization of tourmalines (MR21) and first quartz stage: Aqueous fluids (Lw) trapped at temperatures of 320-470°C and lithostatic pressures (180-280 MPa), gradients of 40-50°C/km

2) Pressure drop >> trapping of Lc and Lc-(w) inclusions

3) Lc-w water-carbonic fluid inclusions from MR19 (quartz lodes) and Lc-w, Lw-c from MR21 (tur quartz) = 240 to 290°C under lithostatic pressures.

Most fluids are predominantly aqueous, with moderate to low salt content and a minor volatile phase of decreasing density

>> composition of the fluids are not representative of demixed fluids from granitic magmas.

>>> fluids linked to migmatisation at depth, interaction with the metamorphic host rocks.





## Main Conclusions

#### Expression of Tourmaline at Mata da Rainha

- Tourmalinisation around syn-deformational quartz veins
- Tourmaline in mineralized quartz veins from the deposit
- Tourmaline in and at the contact with greisens
- Tourmalinized schists cemented by quartz

#### Chemistry of Tourmaline.



- Variations mostly for Fe, Ti, (Mg, Mn and sometime Ca).
- Ti, Fe and Ca more concentrated in the outer dark bands, Mg enriched in the core of zoned tourmalines
- Zn and Mn remain constant (800 to 1000 ppm). Other elements show concentrations of 100 ppm or less

#### Boron isotopic composition of tourmalines

- $\delta^{11}$ B around -10/-11 ‰ are close to those of magmatic tourmalines and S-type granites.
- Very strong tourmalinisation of the schists produces tourmalines whose isotopic characteristics are slightly more extended ( $\delta^{11}B: -7 \text{ to } -13 \%$ ) but centered on the same values.

## P-T Conditions and origin of fluids



- Evolution in three stages, from 320-470°C and lithostatic pressures , 180-280 MPa, gradients of 40-50°C/km to 240 and 100-150 MPa
- Interaction of fluids with the metamorphic host rocks.
- Composition of the fluids are not representative of demixed fluids from granitic magmas

#### Significant mobility of boron by hydrothermal fluids in this area



