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Geochemical and mineralogical features of the Beiras Group metasediments; are they relevant to regional metal endowment?

Ivo Martins; António Mateus M. Helena Hollanda; Colombo Tassinari; Ícaro Dias da Silva; Alcides Pereira



Beiras Group (major elements)

• Complemented with 169 samples collected in the Panasqueira area (Gonçalves et al., 2018);



Beiras Group (minor and trace elements)

- Most significant: Ba, F, Zr, V, Rb, Cr, Zn, B, Li, Ce, Sr, Ni, Cu, Y, La, Ga, Sc, Pb, Nb and Th;
- Positive anomalies (up to ca. 3×AS): Li, Cs, Sn, Hf, Bi, As and Sc;
- Negative anomalies: Mg, Mn, Ca, S, F, Sr, Nb, Ta, W, Th, U, Y, Ni, Mo, Cd, Cu, Pb, Ge and Ag



• High compositional maturity



- High compositional maturity
- Intensely chemical weathered sources (median CIA value 75)



90%

 AI_2O_3

90

100 Kaolinite, Chlorite

10

100

K₂O



R₀_PAN △ MT₀_PAN + LC_ES ♦ R_ES ▲ MT_ES •

- High compositional maturity
- Intensely chemical weathered sources (median CIA value 75)
- Felsic to intermediate igneous sources





- High compositional maturity
- Intensely chemical weathered sources (median CIA value 75)
- Felsic to intermediate igneous sources
- Typical composition of continental island arc sediments





A subset of 30 samples was selected for multisystem (Pb-Sr-Nd) whole rock isotopic analysis



- Variable Rb and Sr contents, although similar order of magnitude.
- Minor Rb/Sr variations; medians range from:
 - 1.82 (lower member) to 2.07 (upper member) in the Malpica do Tejo *Fm*, and

0.8

- 1.29 (distal and lower member) to 1.96 (upper member) in the Rosmaninhal *Fm*.
- Median ⁸⁷Sr/⁸⁶Sr ratios (0.745- 0.751 Malpica do Tejo *Fm*; 0.736-0.751 Rosmaninhal *Fm*) comparable to sedimentary successions in Spain (e.g., Villaseca et al., 2014).
- Median (⁸⁷Sr/⁸⁶Sr)_{560 Ma} values: 0.703 Malpica do Tejo *Fm*; 0.706 to 0.711 Rosmaninhal *Fm*.
- Arguable provenance constraints (Rb/Sr >> 1); post-depositional changes able to affect the concentration of relatively mobile elements.





- Similar range of Sm and Nd for the different members of Malpica do Tejo and Rosmaninhal *Fms*: from ca. 4 to 9 ppm and 20 to 42 ppm, respectively.
- Narrow variations of Sm/Nd (0.20 0.22); median ca. 0.20.
- Median values of εNd range from:
 - -8.654 (lower member) to -7.440 (upper member) in Malpica do Tejo *Fm*;
 - -7.292 (lower member) to -7.582 (upper member) in the Rosmaninhal *Fm*, -7.943 for distal member.
- Median εNd_{560 Ma:} -3.321 and -2.458 for lower and upper members of Malpica do Tejo *Fm;* -1.977 and -2.277 for lower and upper members of Rosmaninhal Fm; -2.689 for distal member of the latter *Fm*. Consistent with data reported in other studies (Pereira et al., 2012; Villaseca et al. 2014)







O Ru □ RI

▲ Rd ● MTu

MTI

× DG

× BG_SF × SN

TDM model ages range from 1.40 to 1.68 Ga. • 5 Median values are: 1.62 and 1.55 Ga for lower and upper 2.5 members of the Malpica do Tejo Fm; 1.48, 1.51 and 1.55 Ga for the lower, upper, ٠ 0 and distal members of the Rosmaninhal Fm. -2.5 Overlap the upper limit of the range reported • е**Nd**_{560Ма} for metapelites forming equivalent units in -5 **Spain** (1.20 to 1.56 Ga, mean of 1.38 Ga; -7.5 Villaseca et al., 2014) **Beiras** Group -10 **Douro Group** 2.5 -12.5 0 "Série Negra" O Ru -2.5 🗆 RI -15 eNd (560Ma) ∆ Rd 1.3 1.35 1.4 1.45 1.5 1.55 1.6 1.65 1.7 1.75 1.8 1.85 1.9 1.95 2 2.05 2.1 2.15 O MTu O MTI TDM (Ga) -7.5 × DG × BG S -10 × SN -12.5

2.05 2.1

1.8 1.85 1.9 1.95 2

TDM (Ga)

1.3 1.35 1.4 1.45 1.5 1.55 1.6 1.65 1.7 1.75

- Considering the bulk Sm-Nd isotopic data, Malpica do Tejo and Rosmaninhal *Fms* were supplied by erosion of crustal rocks, including a noteworthy contribution of felsic (arcderived) igneous sources.
- These igneous sources, younger than 1.2 Ga if the wide range of TDM model ages is considered, may be related to juvenile magmas (e.g., Tassinari et al., 1996) and should document the dismantling of a Cadomian arc located to the south (present coordinates)
- The contribution of felsic (arc-derived) igneous sources is relevant to the metal endowment potentially provided by metasediments when subject to partial melting or intense hydrothermal leaching.





- Similar U and Th contents for all the members of Malpica do Tejo and Rosmaninhal *Fms* ⇒ median U/Th ratios:
 - 0.32 and 0.31 for lower and upper members of Malpica do Tejo *Fm*.
 - 0.30, 0.33 and 0.32 for the lower, upper and distal members of the Rosmaninhal Fm.
- **Pb contents more variable** with median values of:
 - 17.8 and 14.1 ppm for lower and upper members of Malpica do Tejo Fm.
 - 13.7, 14.2 and 20.2 ppm for lower, upper and distal members of the Rosmaninhal Fm
- Isotopic Pb-Pb ratios show minor differences.
- Pb sources fall within the range of crustal, orogen protoliths; signs of a mantle-derived component in some samples from the Malpica do Tejo Fm.
- Pb isotopes consistent with info provided by Sm-Nd systematics.



• Strongly weathered crustal rocks, including significant inputs from felsic (arc-derived) igneous sources

• Loss of Na, Ca and Sr; relative enrichment in K, Rb, Sn, W and elements with similar geochemical behaviour

• Sn and W preferentially incorporated in micas

- Potential to:
 - provide several chemical components when subjected to intense (and long-lasting) hydrothermal leaching;
 - generate Sn- and W-enriched melts if adequate conditions were achieved.

Freixedo do Torrão (Figueira de Castelo Rodrigo)



Experimental determined water-present melting reactions:

- (1) H₂O-saturated *Ms*-granite solidus (Huang and Wyllie, 1973);
- (2) H₂O-saturated solidus in Qz+Or+Ab+H₂O system (Johannes, 1985) with melt aH₂O = 1 (2a), aH₂O = 0.7 (2b), aH₂O = 0.5 (2c), aH₂O = 0.3 (2d) and aH₂O = 0.1 (2e);
- (3) Melting reaction Qz+Pl+Kfs+H₂O with $aH_2O = 1$ (3a) and $aH_2O = 0.1$ (3b) (Stevens and Clemens, 1993);
- (4) Melting reaction Bt+Qz+Kfs+H2O = melt (Peterson and Newton, 1989);
- (5) H₂O-saturated melting of tonalite Bt+Pl+Qz+H₂O = Hbl+melt (Büsch et al., 1974);
- (6) Tonalite H₂O-saturated solidus (Yoder and Tilley, 1962);
- (7) H₂O-saturated melting reaction Qz+Kfs+Als+H₂O = melt (Johannes and Holtz, 1996);
- (8) H₂O-saturated granite solidus (Qz+Ab+Or+H₂O=melt) (Ebadi and Johannes, 1991);
- (9) Melting reaction (Bt+Pl+Qz+Kfs+H₂O=Hbl+Grt+melt) of gneiss with 4wt% H2O added (Gardien et al., 2000).

Experimental determined **dehydration** melting reactions:

- (1) Ms-dehydration melting of *Ms-Bt*-schist (1a) and *Ms*-schist (1b) (Patiño Douce and Harris, 1998);
- (2) Ms-dehydration melting (Pëto, 1976);
- (3) Bt-dehydration melting of metagreywacke (Vielzeuf and Montel, 1994);
- (4) Bt-dehydration melting of metagreywacke (4a) and gneiss(4b) (Patiño Douce and Beard, 1995, 1996);
- (5) Bt-dehydration melting in MASH system (Vielzeuf and Clemens, 1992);
- (6) Bt-dehydration of gneiss (Gradien et a., 2000)
- (7) Bt-dehydration melting of metapelite (Le Breton and Thompson, 1988).

In comparison, note different slopes and solidus T. Waterpresent melting results in negative volume change whereas dehydration melting leads to positive volume change (\Rightarrow fracturing and melt expulsion into the host).



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Thank you for your attention!

Compositionally modified metasediments (Cabeço de Argemela)