



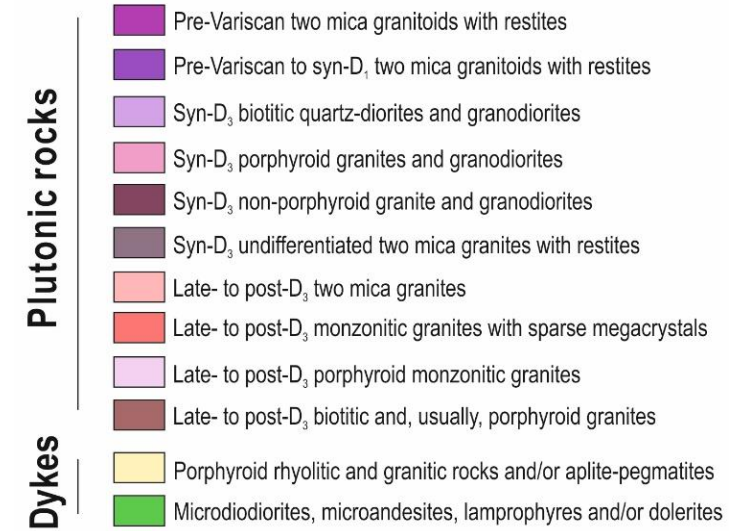
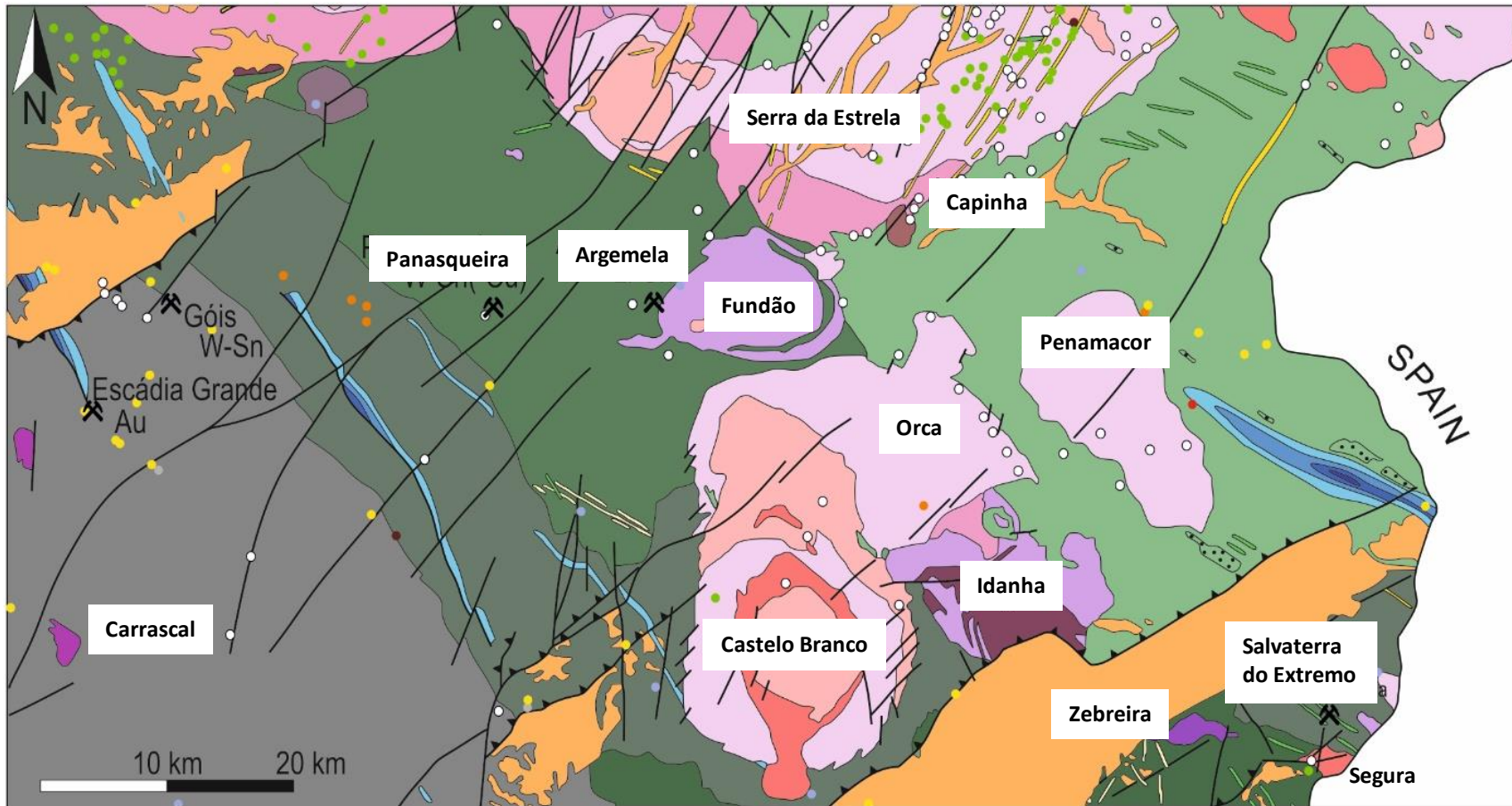
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# Compositional attributes and geochronology of granite suites; implications to metallogenic processes

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# Granitoid, aplite & pegmatite rocks



## However:

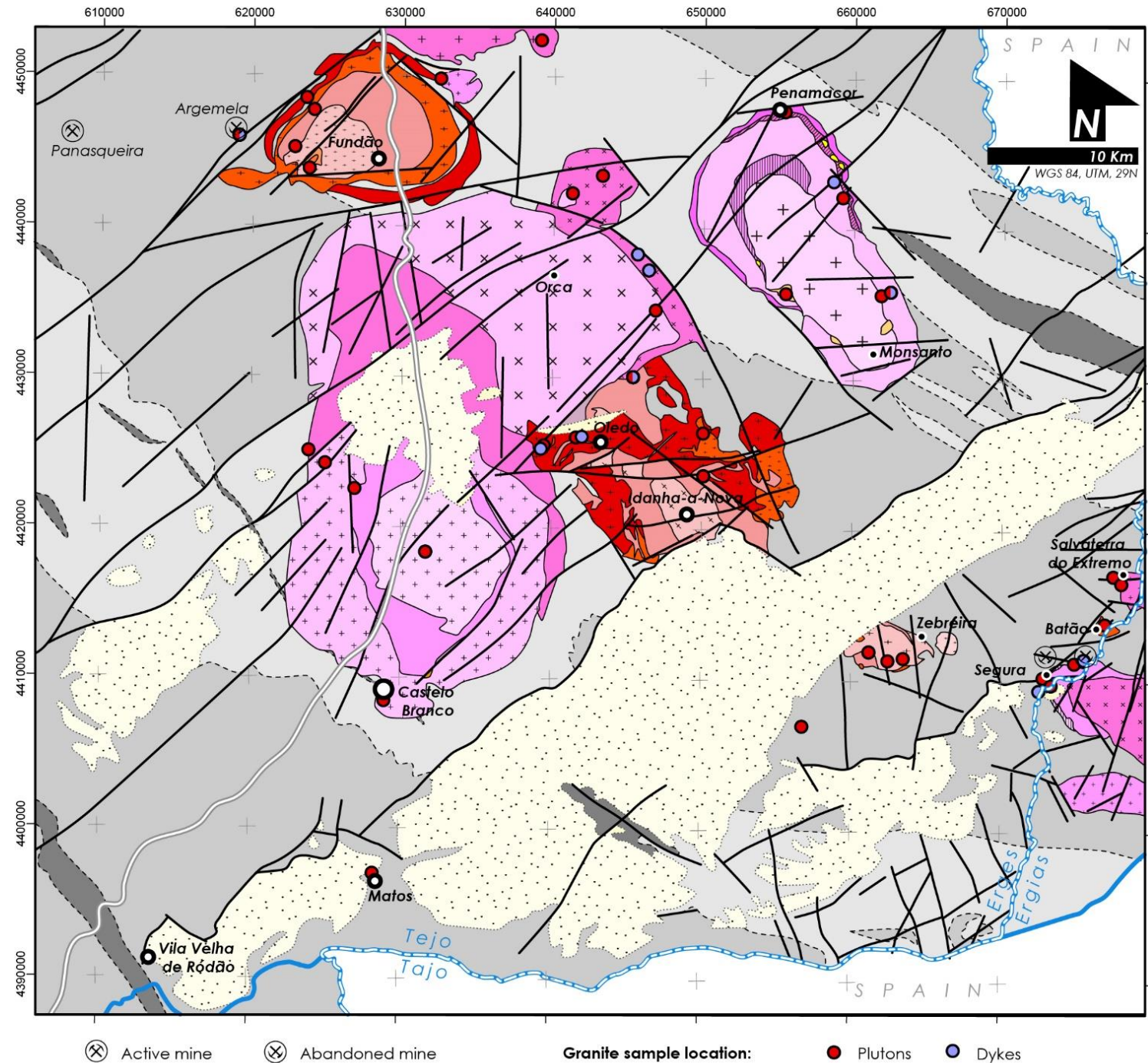
- Some data needing validation;
- Several gaps demanding further research (facies and elements analysed); and
- Limited information for aplite and pegmatite bodies.

Starting point: results from many studies carried out throughout the past four decades



# Sampling

- Focus on the Panasqueira-Segura area.
- 62 samples of granitoid rocks + 45 samples of aplite-pegmatite bodies and porphyry dykes.
- *23 extra samples from Fundão and Orca (Mata da Rainha) recently collected (Sept2023) and under study.*
- Compared with 192 whole-rock analyses compiled from published data.





# Two main suites, documenting distinct magmatic events

- Cambrian-Ordovician**

## Early Ordovician plutons

### Fundão

- G<sub>FUN</sub>1: Biotite monzonitic granite
- G<sub>FUN</sub>2: Biotite quartz-diorite
- G<sub>FUN</sub>3: Biotite granodiorite
- G<sub>FUN</sub>4: Biotite quartz-diorite
- G<sub>FUN</sub>5: Transitional facies granite
- G<sub>FUN</sub>6: Monzonitic granite

### Oledo - Idanha-a-Nova

- G<sub>OIN</sub>1: Medium-grained biotite granodiorite
- G<sub>OIN</sub>2,3: Medium- to coarse-grained porphyritic biotite-muscovite granodiorite
- G<sub>OIN</sub>4: Medium- to coarse-grained muscovite-biotite granite
- G<sub>OIN</sub>5: Muscovite-biotite porphyroid granite

### Zebreira

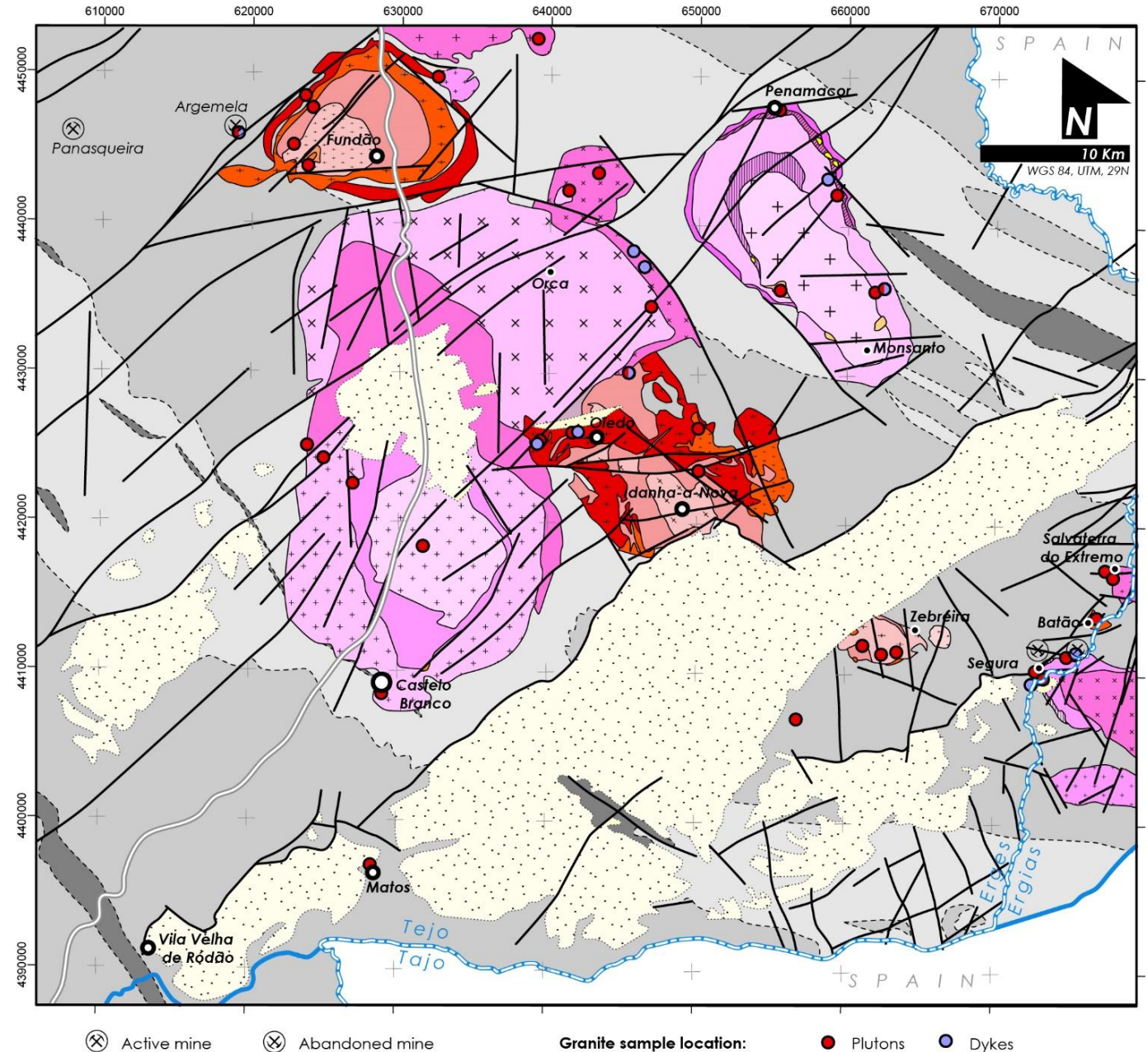
- Medium-coarse grained granodiorite
- Metassomatized granodiorite
- Porphyroid muscovite granite
- Medium-grained monzonitic granite

### Batão

- Tonalite

- Cenozoic cover
- Ordovician-Silurian synclines
- Rosmaninhal Fm. (Ediacaran-lower Cambrian)
- Malpica do Tejo Fm. (Ediacaran)

- Stratigraphic boundary
- Intrusive contact
- Fault
- River
- Spanish-Portuguese border





# Two main suites, documenting distinct magmatic events

- Carboniferous-Permian

**Variscan (late Carboniferous) plutons**

**Pero-Viseu and Capinha**

Coarse- to very coarse-grained, porphyritic granite

**Afalaia**

Albite-muscovite granite

**Orca**

G<sub>ORC</sub> 1: Medium- to fine-grained muscovite granite

G<sub>ORC</sub> 2: Porphyroid biotite-muscovite granite

**Castelo Branco**

G<sub>CB</sub> 1: Medium- to fine-grained muscovite-biotite granite

G<sub>CB</sub> 2: Medium- to fine-grained porphyritic biotite-muscovite granodiorite

G<sub>CB</sub> 3,4: Medium- to coarse-grained porphyritic biotite-muscovite granite

G<sub>CB</sub> 5: Coarse-grained muscovite-biotite granite

**Penamacor-Monsanto**

G<sub>PM</sub> 1: Medium- to coarse-grained muscovite-biotite granite

G<sub>PM</sub> 2: Medium-grained muscovite-biotite granite

G<sub>PM</sub> 3: Coarse- to medium-grained porphyritic biotite-muscovite granite

G<sub>PM</sub> 4: Medium-grained porphyritic biotite-muscovite granite

G<sub>PM</sub> 5: Coarse-grained porphyritic muscovite-biotite granite

G<sub>PM</sub> 6: Medium- to coarse-grained porphyritic

Pegmatite-Aplite dykes swarm

**Salvaterra do Extremo**

Medium- to coarse-grained muscovite granite

**Segura-Cabeza de Araya**

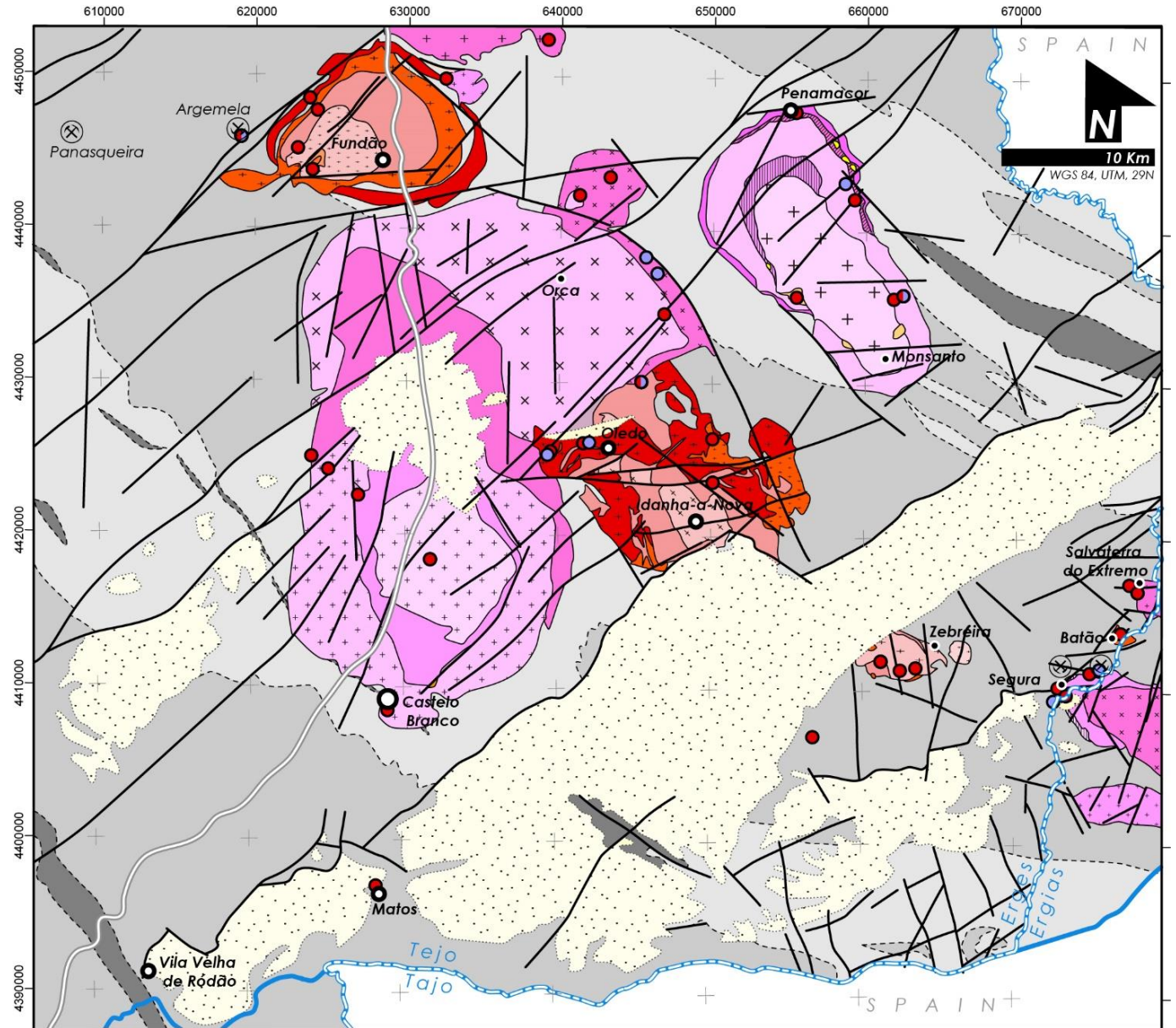
G<sub>SCA</sub> 1: Coarse- to fine-grained muscovite-turmaline granite

G<sub>SCA</sub> 2: Porphyroid two-mica granite

G<sub>SCA</sub> 3: Porphyroid biotite-cordierite granite

**Estorniões**

Two-mica porphyroid granite



(X) Active mine

(X) Abandoned mine

Granite sample location:

● Plutons

● Dykes

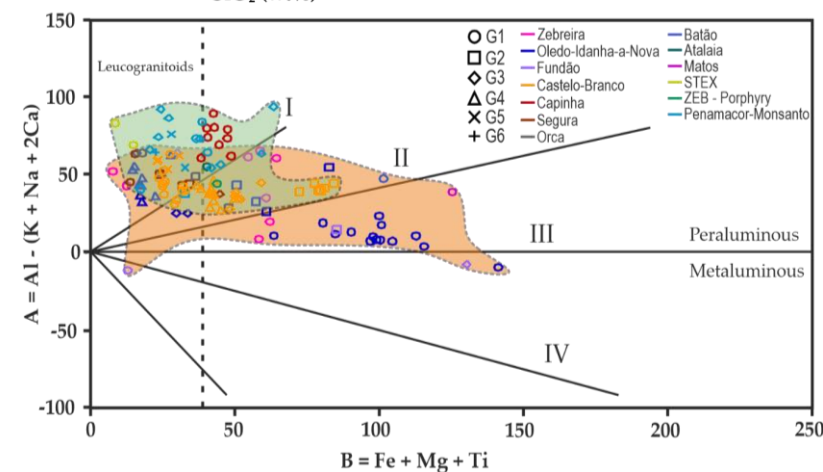
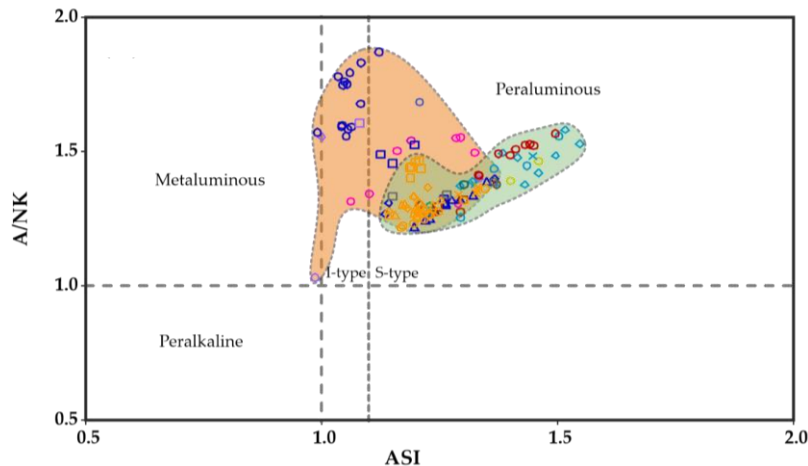
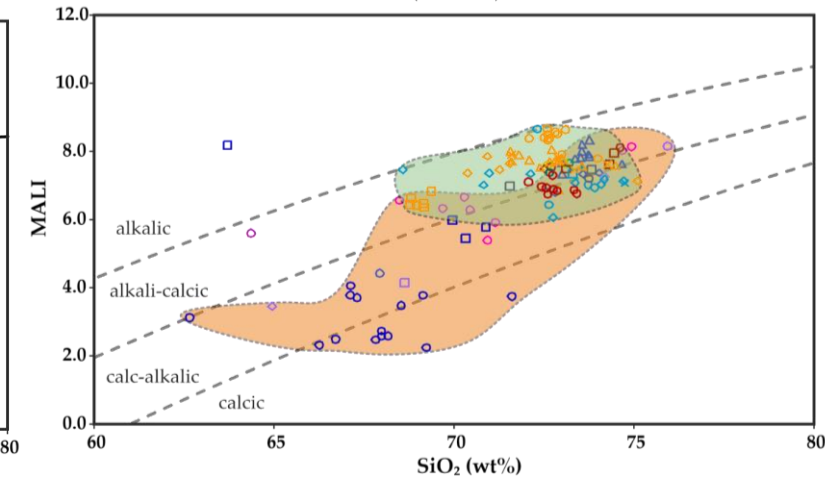
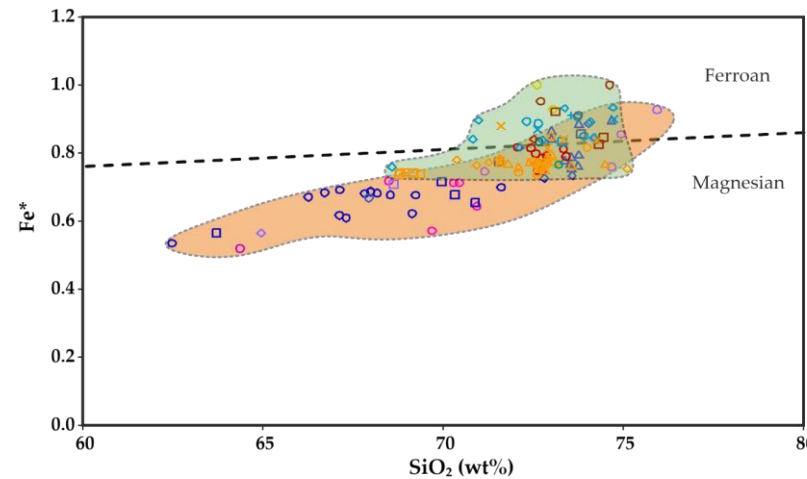
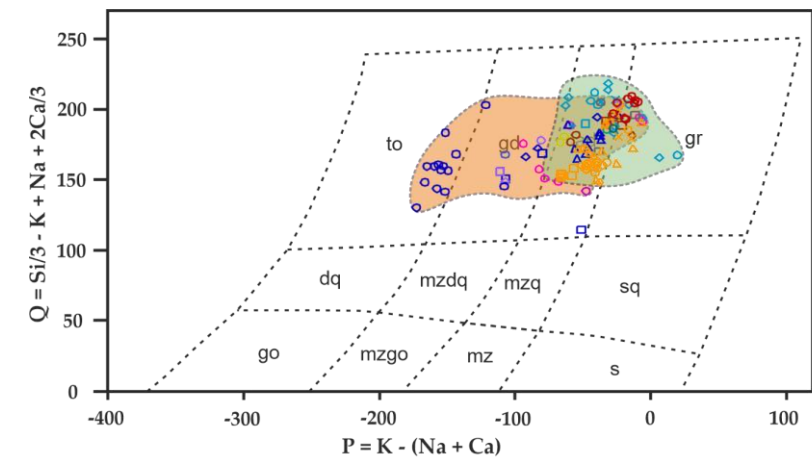
# Compositional features

## - Cambrian-Ordovician magmatism:

- Weakly peraluminous I-type;
- Bt/Bt>Ms tonalite to granodiorite;
- Calcic to calc-alkalic series and magnesian granitoid rocks

## - Variscan magmatism:

- Highly peraluminous S-type;
- Ms>Bt/Bt>Ms monzogranite to granite;
- Calc-alkalic to alkali-calcic and magnesian to ferroan granites





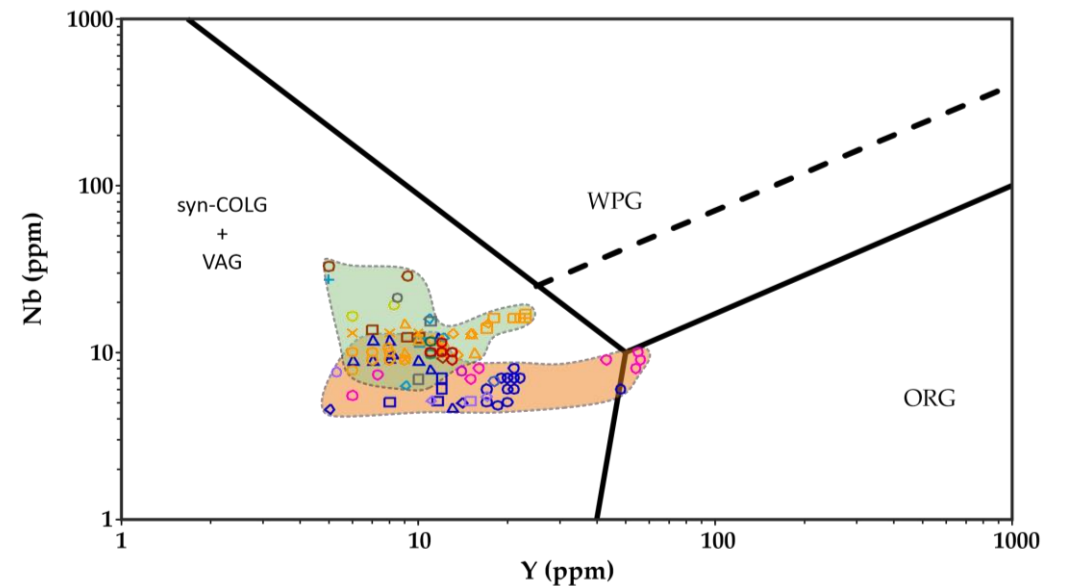
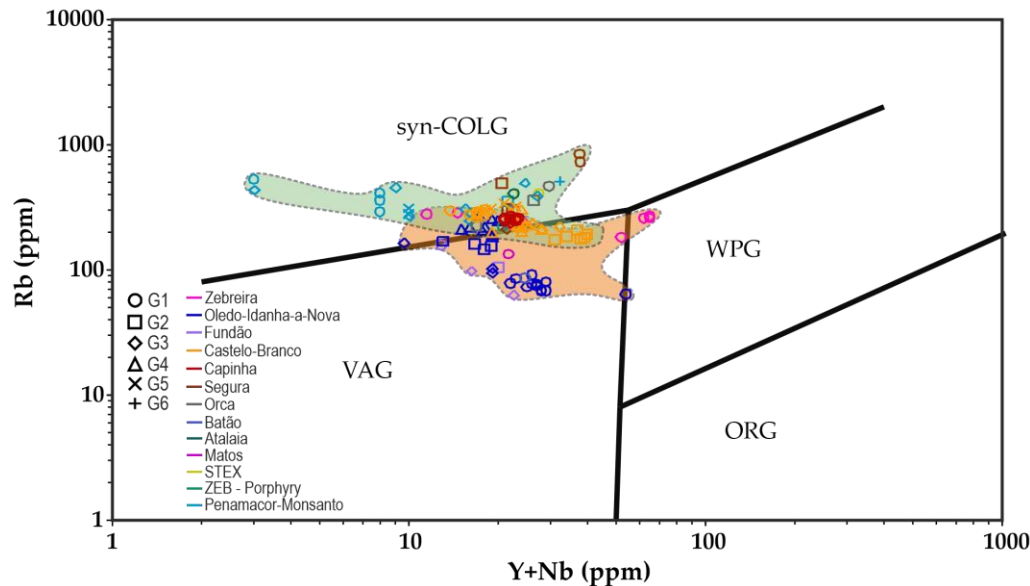
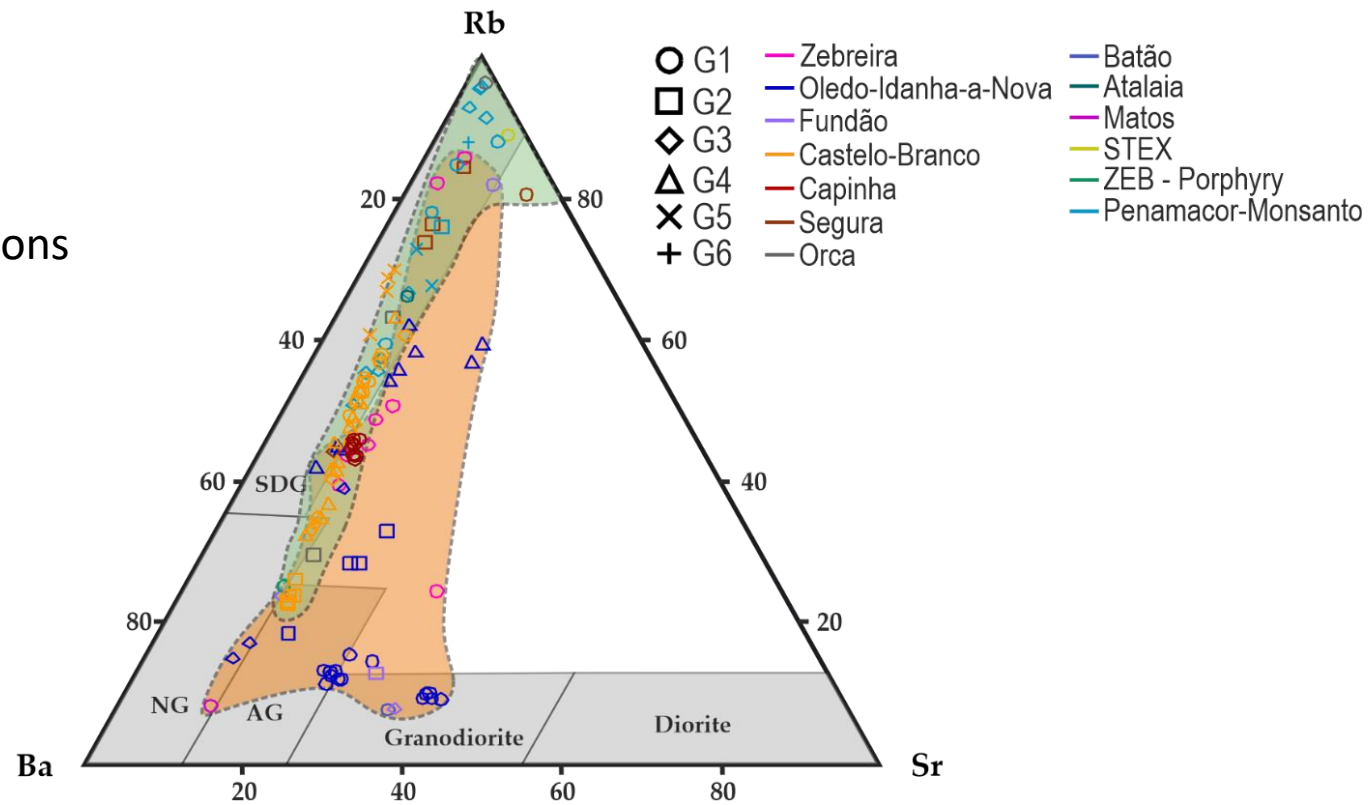
# Compositional features

## - Cambrian-Ordovician magmatism:

- Tonalite/granodiorite to normal granite compositions (condensed differentiation trends);
- Volcanic arc granitoids.

## - Variscan magmatism:

- Strongly differentiated rocks;
- Syn-collisional granites.



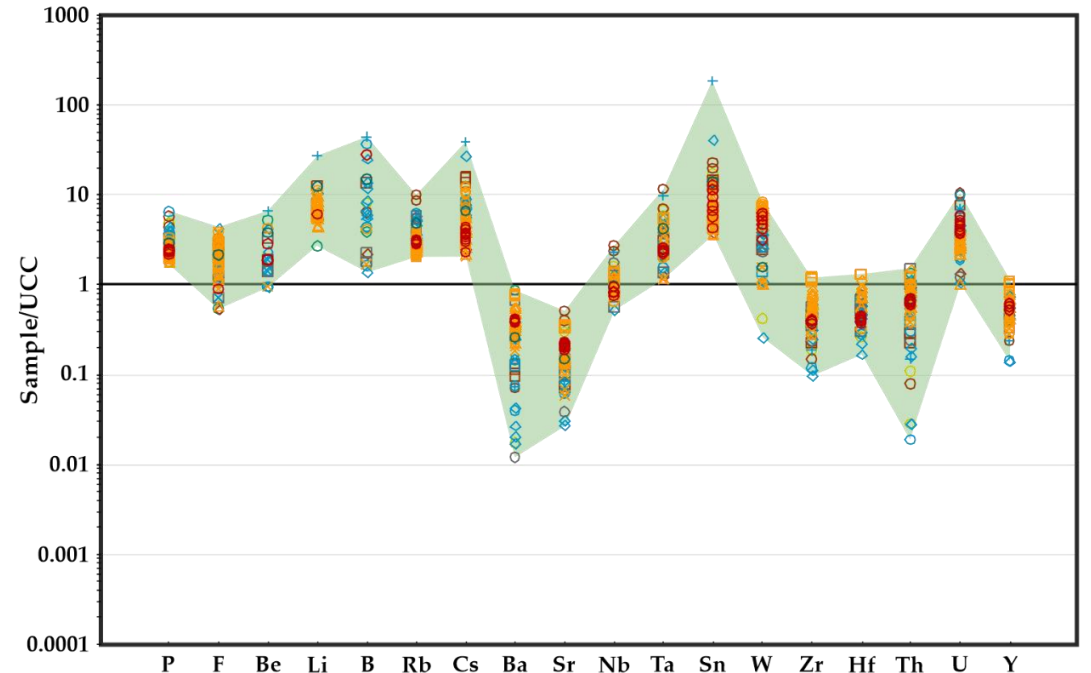
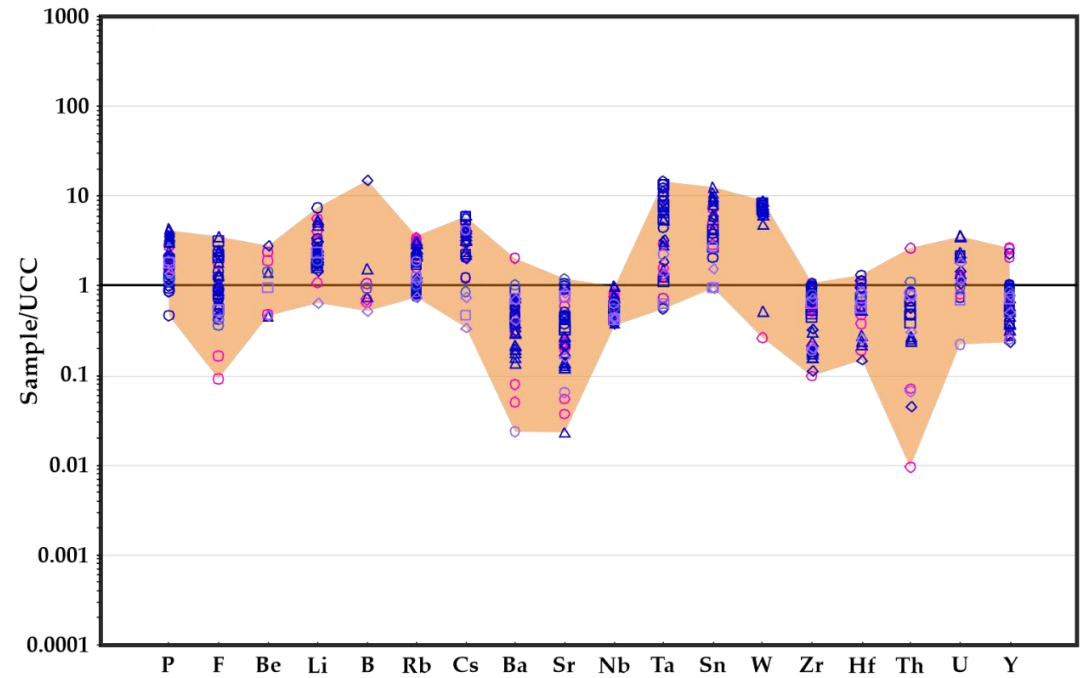
# Compositional features

## - Cambrian-Ordovician magmatism:

- Enrichment (up to 20×UCC) in P, F, Be, Li, B, Rb, Cs, Ta, Sn, W and U;
- Depletion in Ba, Sr, Zr, Hf, Th and Y.

## - Variscan magmatism:

- Prominent enrichments (up to 200×UCC, usually ranging from 20 to 50×UCC) and depletions in similar elements





# Compositional features | Whole rock multi-system isotope analysis

## Cambrian-Ordovician magmatism (11 samples)

- Main granitoid facies, excluding those from Idanha-a-Nova – Oledo (already reported in Antunes et al. 2010).
- Two-mica granite from the latter pluton and some highly differentiated facies, mostly forming dykes criss-crossing the granitoid facies (some of them also dated as Early Ordovician).

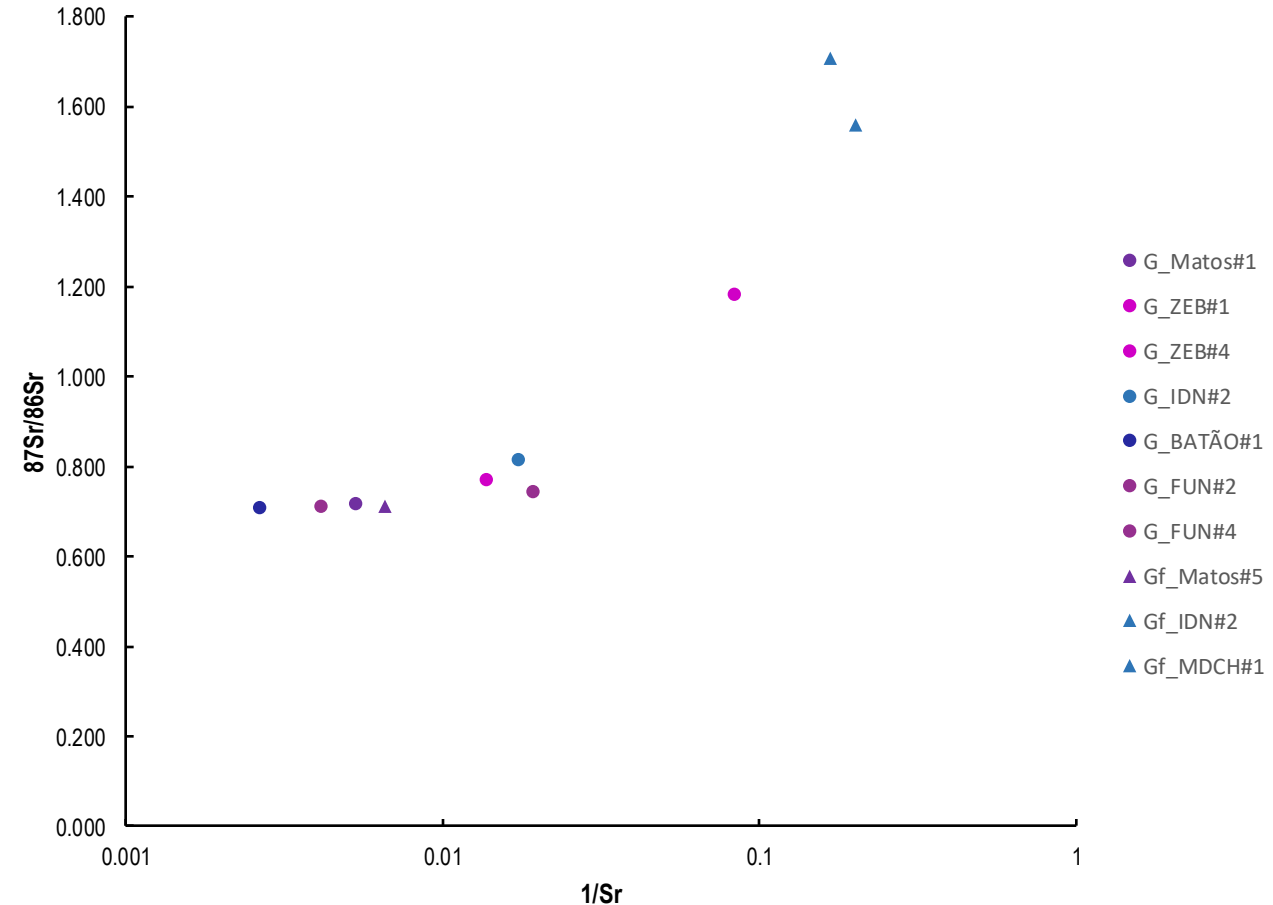
## Variscan magmatism (22 samples)

- Main granite facies and dyke arrays from Salvaterra do Extremo, Segura, Penamacor-Monsanto (including the differentiated *ms*-rich facies of Medelim), Orca and Argemela, and excluding those from Castelo Branco (reported in Antunes et al. 2008).
- Several other facies, namely:
  - The *bt*-rich granite body of Capinha, to the north of Orca and Penamacor-Monsanto plutons;
  - The hidden two-mica granite of Panasqueira (SCB2#11) and an aplitic dyke (SMN1#4), both intersected by drilling, besides the greisen nearby the underground mined quartz lodes (P11-G);
  - A late granite body criss-crossing the Fundão granitoids (G-LP#1); and
  - The late porphyry rocks forming laccoliths and dykes at Marcelina and Furão (Zebreira area).

# Compositional features | Cambrian-Ordovician magmatism

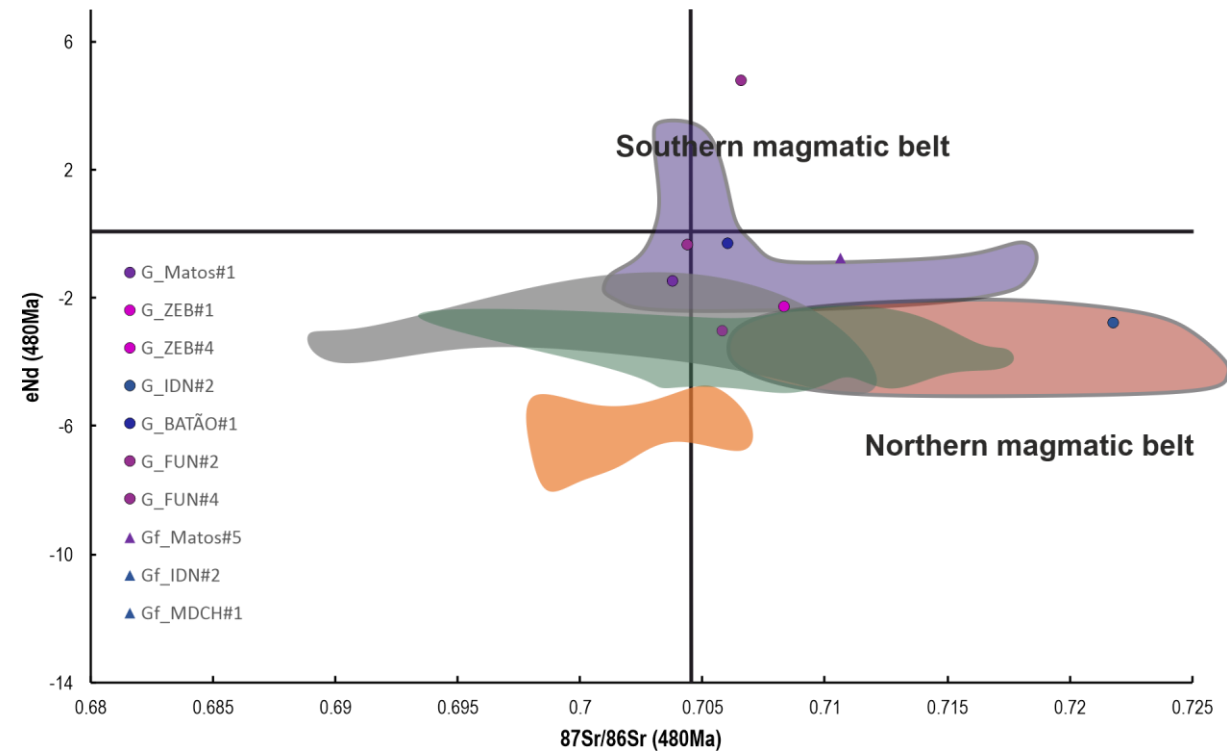
## Sr systematics

- Variable Sr and Rb contents, causing large Rb/Sr fluctuations:
  - 0.23 in the Batão tonalite;
  - 3.16 - 23.17 in granitoid facies from Zebreira;
  - 0.43 - 1.90 in granitoid facies from Fundão;
  - 4.71 in the two-mica granite, and up to 73.50 in highly differentiated dykes from Idanha-a-Nova – Oledo; and
  - 0.71 in granitoids from Matos.
- $^{87}\text{Sr}/^{86}\text{Sr}_{480 \text{ Ma}}$  values are:
  - 0.706 for Batão,
  - 0.704-0.709 for Fundão,
  - 0.704-0.711 for Matos,
  - 0.707-0.708 for Zebreira, and
  - 0.722 for the Idanha-a-Nova – Oledo leucogranite.
  - **Comparable with ranges reported for similar rocks across the CIZ southern domain and other Variscan orogenic segments (e.g., Talavera et al., 2013).**



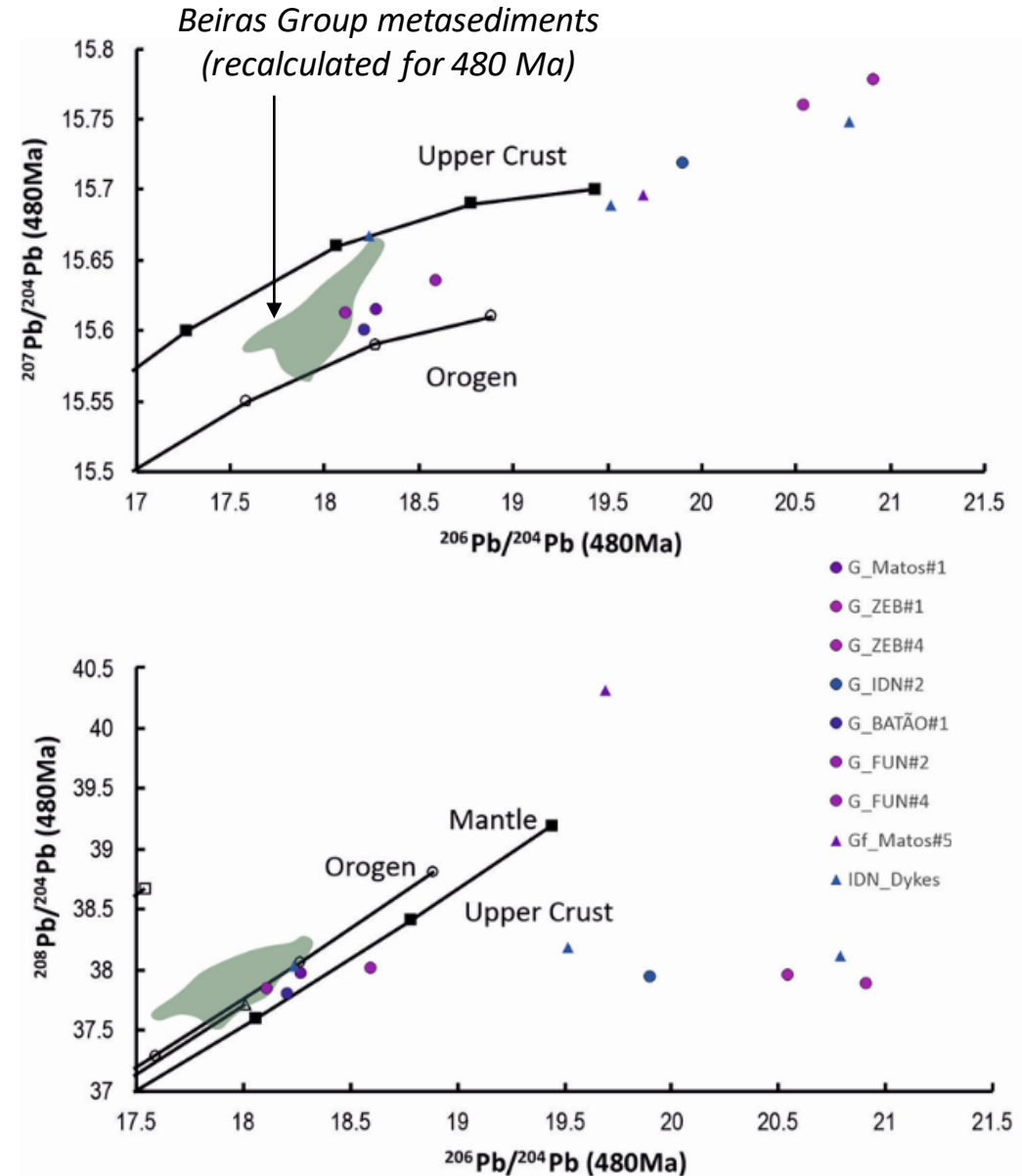


- **Limited variation of Sm and Nd contents, and Sm/Nd ratios (0.18-0.28).**
- **Positive  $\epsilon\text{Nd}_{480 \text{ Ma}}$**  for one of the granitoid facies forming the Fundão pluton (4.783; G-FUN#4) and for two of the dykes from Idanha-a-Nova – Oledo (0.756 and 1.008).
- **$-0.331 \leq \epsilon\text{Nd}_{480 \text{ Ma}} \leq -1.485$**  for granitoid rocks from Matos, Batão and Fundão.
- **$-3.048 \leq \epsilon\text{Nd}_{480 \text{ Ma}} \leq -2.788$**  for granitoids from Zebreira and the two-mica granite from Idanha-a-Nova – Oledo.
- **TDM model ages define two main groups:**
  - The older, comprises the two-mica granite from Idanha-a-Nova – Oledo (ca.1.50 Ga) and the granitoids from Zebreira (ca. 1.40 to 1.50 Ga);
  - The younger, includes the Batão tonalite (1.24 Ga), the granitoids from Matos (1.27 to 1.33 Ga), the dykes from Idanha-a-Nova – Oledo (1.15 to 1.33 Ga) and the granitoids from Fundão (0.82 to 1.24 Ga).



*Compositional fields of meta-igneous rocks from the two C-O magmatic belts (purple and dark-pink areas) and the metasedimentary Neoproterozoic succession in each subzone (grey = Série Negra; orange = Douro Group; green = Beiras Group). Details on the isotopic data in Villaseca et al. (2014) and references therein.*

- Variable U (2-10 ppm), Th (0.8-28 ppm), and Pb (7-30 ppm) contents.
- Usually,  $Th > U$  and  $0.20 < U/Th < 0.38$ ; exceptions are the *bt*-rich granites from Zebreira and Fundão (G-ZEB#4 and G-FUN#4), the two-mica granite from Idanha-a-Nova – Oledo (G-IDN#2) and two dykes from the latter pluton (Gf-IDN#2 and Gf-MDC#1), which show  $1.5 \leq U/Th \leq 23.3$ .
- Pb-Pb isotopic ratios (480 Ma) spreading between the “*orogen*” and “*upper crustal*” reference curves.
- Slight overlapping with the Beiras Group metapelites  $\Rightarrow$  protoliths isotopically distinct from these metasediments involved in crustal melting processes. Main deviations are granitoids from Zebreira and the two-mica granite from Idanha-a-Nova – Oledo, all displaying negative  $\epsilon Nd_{480 \text{ Ma}}$  values and older TDM model ages.





## Cambrian-Ordovician magmatism

*(Batão, Zebreira, Idanha-a-Nova-Oledo, Fundão, Matos)*

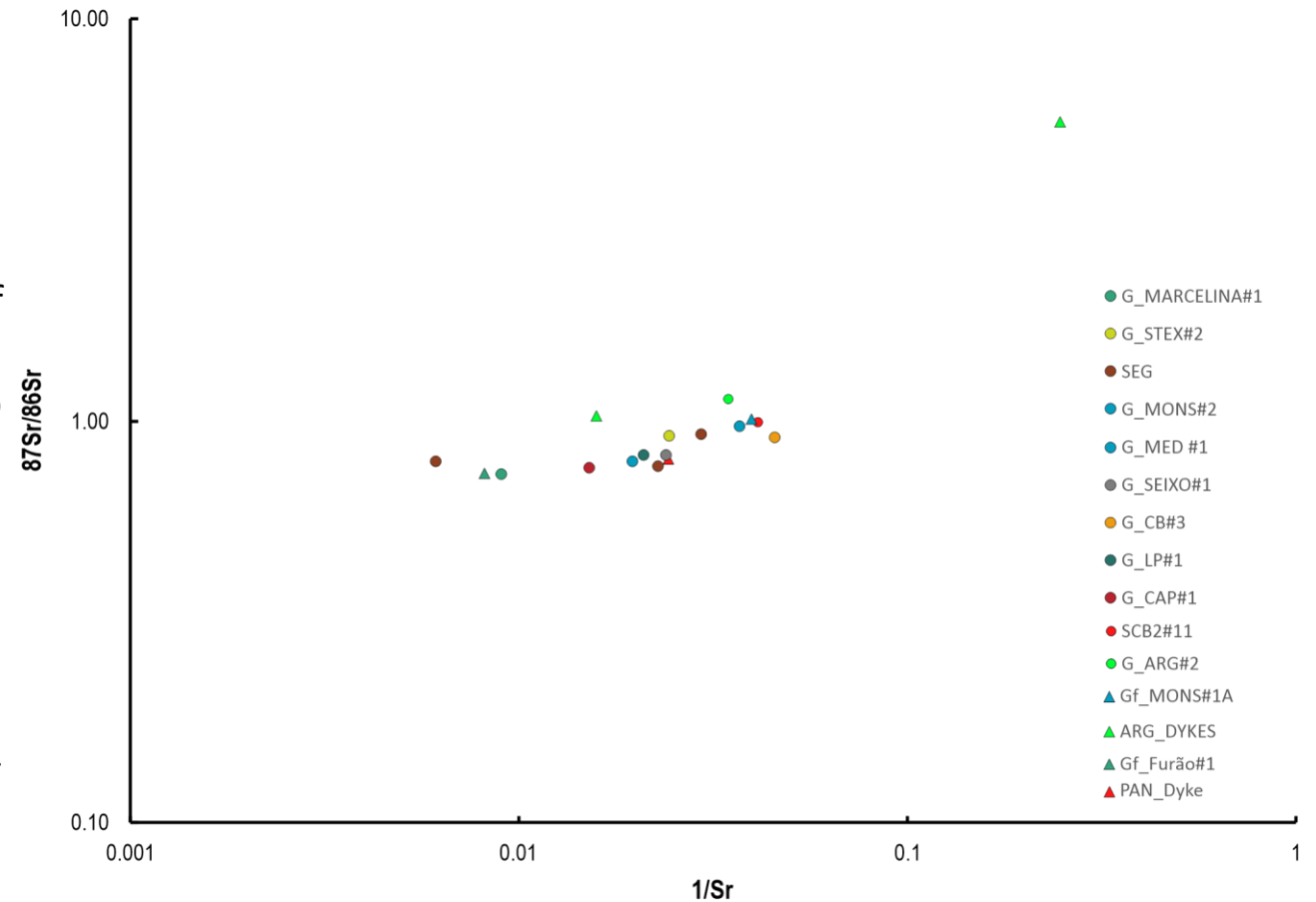
- Independent magma batches reflecting different proportions of mixing between melts that preserve a “*orogen*” Pb-signature with melts generated in upper crustal levels (with variable degrees of partial melting).
- Consistent with the proposal of Castro et al. (2020) for the formation of granodiorite/tonalite plutons and minor intrusions during Cambrian-Ordovician times:
  - Emplacement of sanukitoid magmas at the lower crust, transferring the heat and releasing water during their crystallisation;
  - Subsequent partial melting of the lower crust, which includes the roots of the calc-alkaline Cadomian arc; and
  - Further magma rising through the crust, assimilating and/or triggering partial melting of metasediments (deeper portions of the siliciclastic succession forming the Beiras Group).



# Compositional features | Variscan magmatism

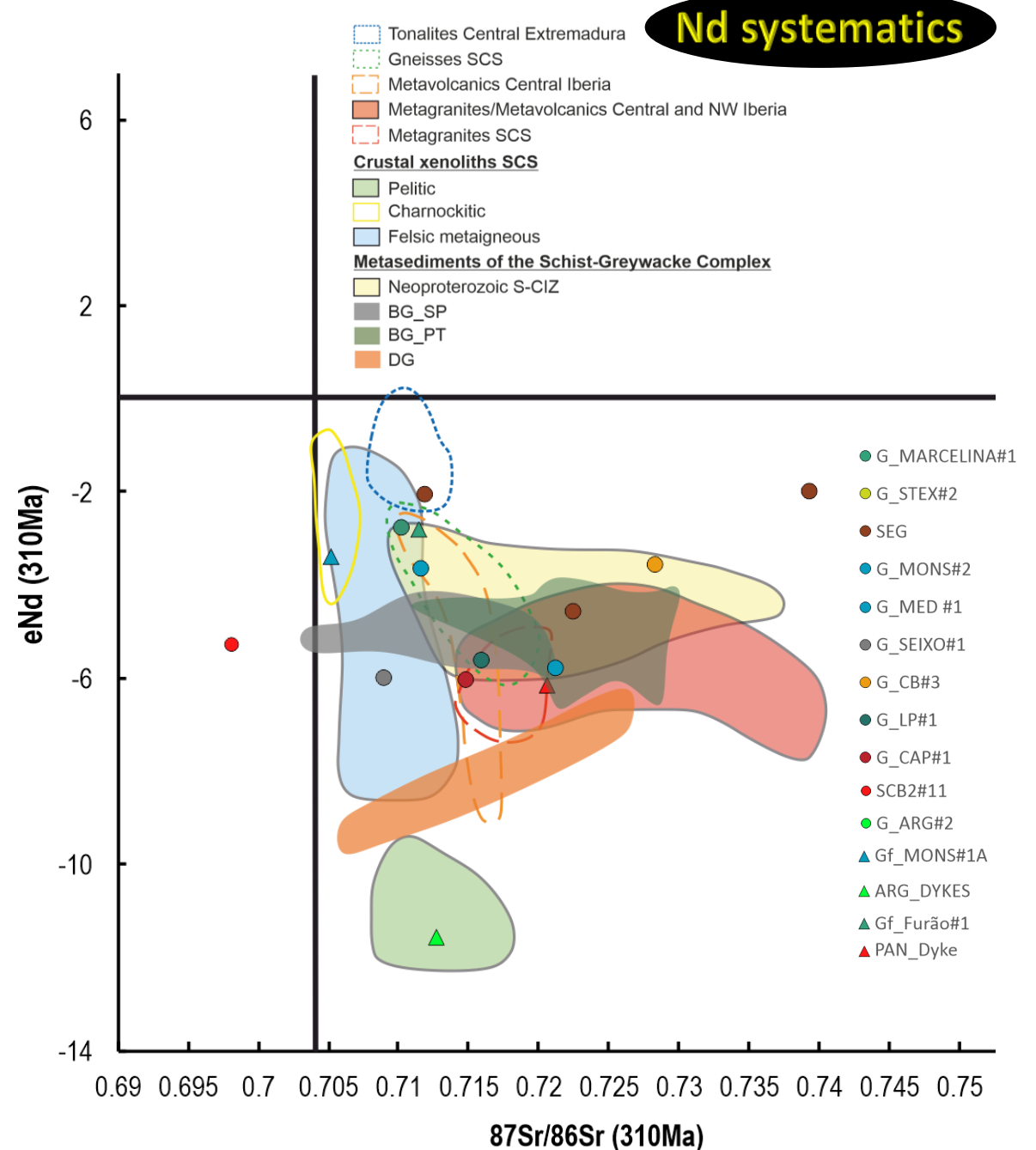
## Sr systematics

- **Wide range** of Rb (223-845 ppm) and Sr (27-164 ppm) contents; remarkable **Rb enrichment in the Argemela suite** (1110-1480 ppm) + lower Sr values (4-29 ppm)
- **Usually,  $2 \leq \text{Rb/Sr} \leq 10$** , but rise to
  - 14-20 in highly differentiated facies of Segura, Medelim and Castelo Branco.
  - 23-30 in pegmatites (e.g. Monsanto) and granitic facies recording compositional changes concurrent of mineralising processes (e.g. Panasqueira)
  - spread up to 232 in cases of exceptional Rb enrichments (Argemela).
- $^{87}\text{Sr}/^{86}\text{Sr}_{310 \text{ Ma}}$ : **0.705-0.716** in “common” granites; **0.721-0.729** in granites showing higher differentiation degrees and/or post-emplacement compositional changes; **0.710-0.712** in late porphyry rocks.
  - First subset compares well with “late to post- $D_3$ , *bt*-dominant Iberian Variscan suites” ( $\approx 312$ -300 Ma,  $0.7064 \leq ^{87}\text{Sr}/^{86}\text{Sr} \leq 0.7085$ ).
  - The other two subsets deviate significantly from the “signature” ascribed to the post- $D_3$  ( $\approx 299$ -290 Ma) granite suite ( $0.7033 \leq (^{87}\text{Sr}/^{86}\text{Sr})_i \leq 0.7079$ ).



# Compositional features | Variscan magmatism

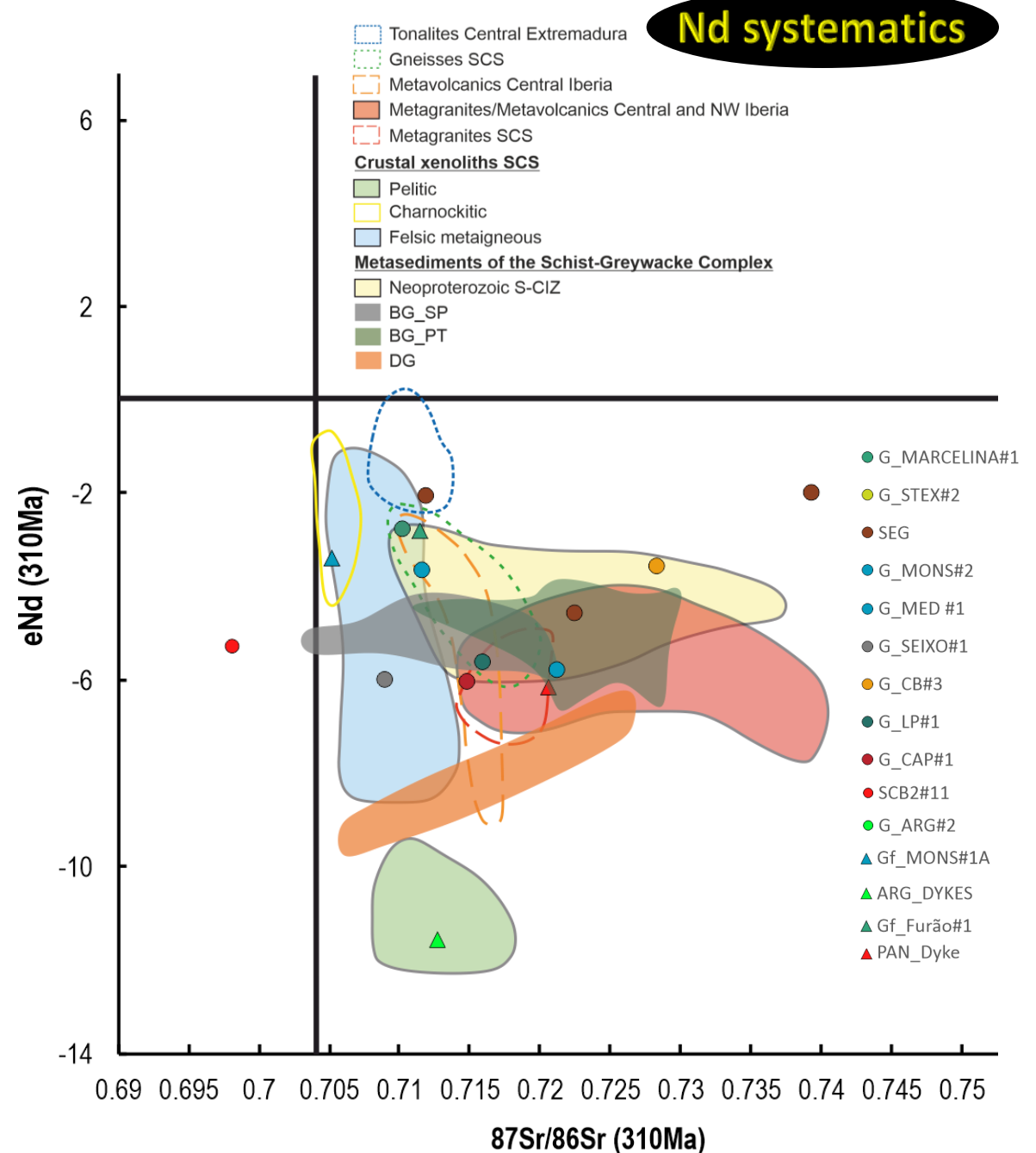
- **Variable Sm (0.7-3.8 ppm ) and Nd (1.3-18.7 ppm) contents.** Argemela samples + Panasqueira aplitic dyke show lower Sm (0.05-0.43 ppm) and Nd (0.13-1.05 ppm), also diverging from the ranges ( $3.9 \leq \text{Sm} \leq 4.5 \text{ ppm}$ ;  $24 \leq \text{Nd} \leq 27 \text{ ppm}$ ) in late porphyry rocks.
- **Sm/Nd ratio confined to the 0.20-0.54 range; 0.16-0.17 in late porphyry rocks.**
- $\epsilon\text{Nd}_{310 \text{ Ma}}$  vary between:
  - ca. -6.5 and -5.7 in granites from Capinha, Salvaterra do Extremo, Medelim, Orca and Fundão, and aplite dykes from Argemela and Panasqueira;
  - ca. -5 and -3 in the inner granite facies of Segura, granite and pegmatite of Monsanto, Castelo Branco (differentiated granite), Panasqueira (two-mica granite) and Argemela (differentiated/modified granite);
  - ca. -2.9 and -2.7 in late porphyry rocks;
- $\epsilon\text{Nd}_{310 \text{ Ma}}$  is around:
  - -2 for the *ms*-rich granite border facies of Segura; and
  - -11.6 for one of the aplite dykes sampled in Argemela



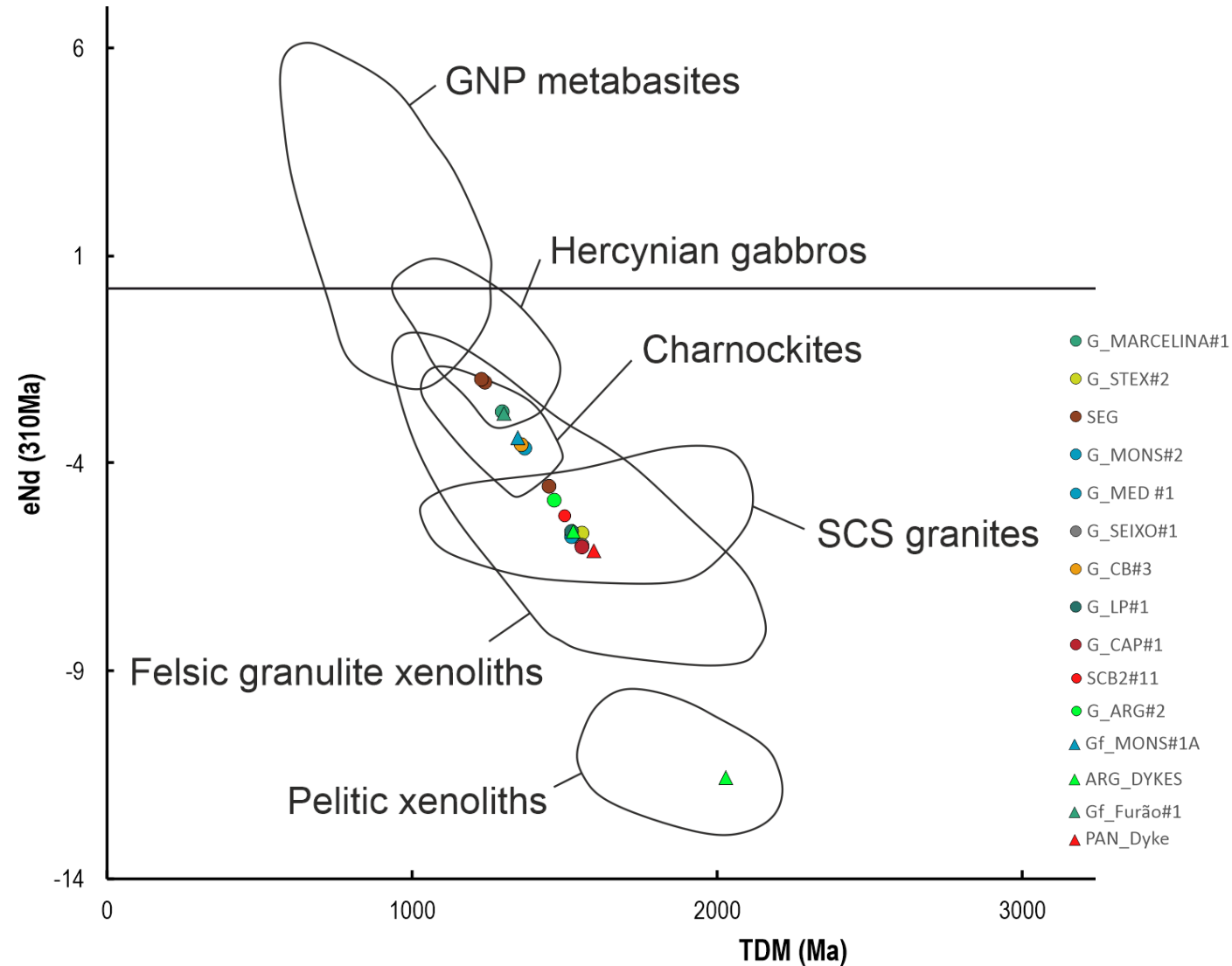


# Compositional features | Variscan magmatism

- $-6.5 \leq \epsilon Nd_{310 \text{ Ma}} \leq -5.7 \equiv$  ranges of “late to post- $D_3$ , *bt*-dominant suites” ( $-6.2 \leq \epsilon Nd_i \leq -4$ ) in CIZ (e.g., Ribeiro et al., 2019). Prevalent crustal origin of the melts also consistent with  $^{87}\text{Sr}/^{86}\text{Sr}_{310 \text{ Ma}}$  values.
- **Moderate negative  $\epsilon Nd_{310 \text{ Ma}}$  values  $\Rightarrow$  other protoliths than metapelites in partial melting processes** (meta-greywackes and/or meta-igneous sources).
- $\epsilon Nd_{310 \text{ Ma}}$  shifting (-5 to -3)  $\Rightarrow$  higher heterogeneity of the crustal source region + possible contributions of deeper sources (inc. lower crust partial melting).
- **Late porphyry rocks ( $\epsilon Nd_{310 \text{ Ma}} \approx -2.8$ ) = melting of lower crustal sources**, although contaminated with pelitic-derived components during their ascent in the crust ( $^{87}\text{Sr}/^{86}\text{Sr}_{310 \text{ Ma}} \approx 0.711$ ).
- **Aplite dyke in Argemela (lowest  $\epsilon Nd_{310 \text{ Ma}}$ ; -11.6) = partial melting of middle to lower crustal sources with dominant pelitic composition** ( $^{87}\text{Sr}/^{86}\text{Sr}_{310 \text{ Ma}} = 0.713$ ).



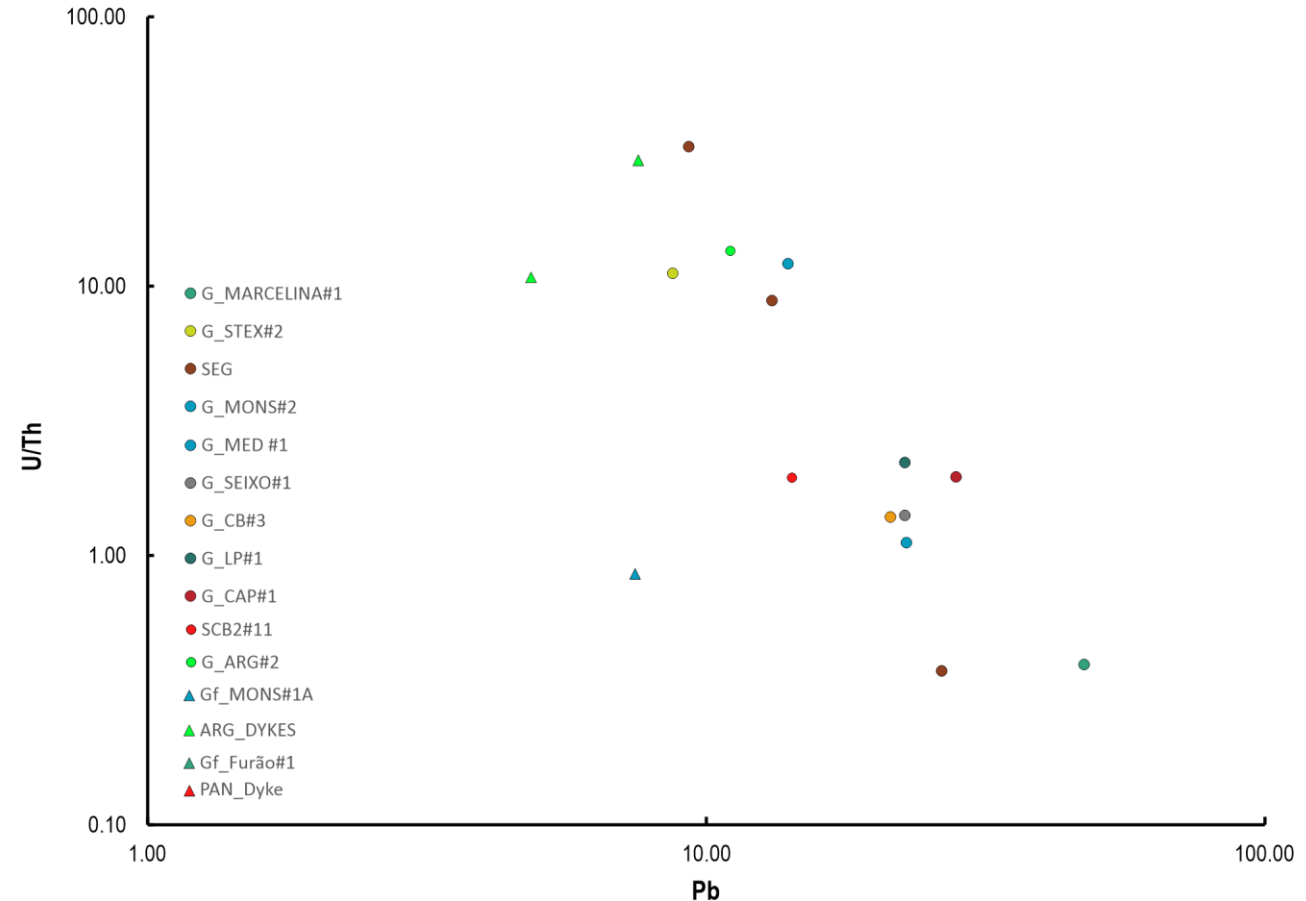
- Inferences compatible with the estimated TDM model ages:
  - **ca. 2 Ga** for the Argemela aplite dyke with the lowest  $\epsilon\text{Nd}_{310\text{ Ma}}$  value;
  - **ca. 1.6 to 1.5 Ga** for the Capinha, Salvaterra do Extremo, Medelim, Orca, Fundão and Panasqueira granites, as well as for the aplite dykes from Argemela and Panasqueira;
  - **ca. 1.4 to 1.3 Ga** for samples from Segura (granite inner facies), Monsanto (granite and pegmatite), Castelo Branco (differentiated granite), and Argemela (differentiated/modified granite);
  - **ca. 1.3 Ga** for the late porphyry rocks; and
  - **ca. 1.2 Ga** for the *ms*-rich granite border facies of Segura.



# Compositional features | Variscan magmatism

## Pb systematics

- **$10 \leq \text{Pb} \leq 25 \text{ ppm}$** , but
  - $\approx 28 \text{ ppm}$  for Capinha granite,  $\approx 48 \text{ ppm}$  for Marcelina porphyry, and  $123 \text{ ppm}$  for Panasqueira greisen.  $\text{Pb} < 10 \text{ ppm}$  for Salvaterra do Extremo granite, one granite facies of Segura, Monsanto pegmatite and all the dykes sampled at Segura and Argemela.
- **$\approx 11 \leq \text{U} \leq 27 \text{ ppm}$** , although
  - $\approx 3\text{--}8 \text{ ppm}$  for some granitic facies of Segura, Castelo Branco and Argemela, as well as the Monsanto pegmatite and the Marcelina porphyry.
- **$\approx 1 \leq \text{Th} \leq 11 \text{ ppm}$** , reaching  $14 \text{ ppm}$  in Marcelina porphyry.
- **Based on the U/Th ratio, 4 subgroups could be distinguished**, but mostly importantly **higher U/Th ratios ( $> 10$ ) characterise granite and dyke facies showing signs of mineralisation and/or effects of mineral/textural changes triggered by mineralising processes.**

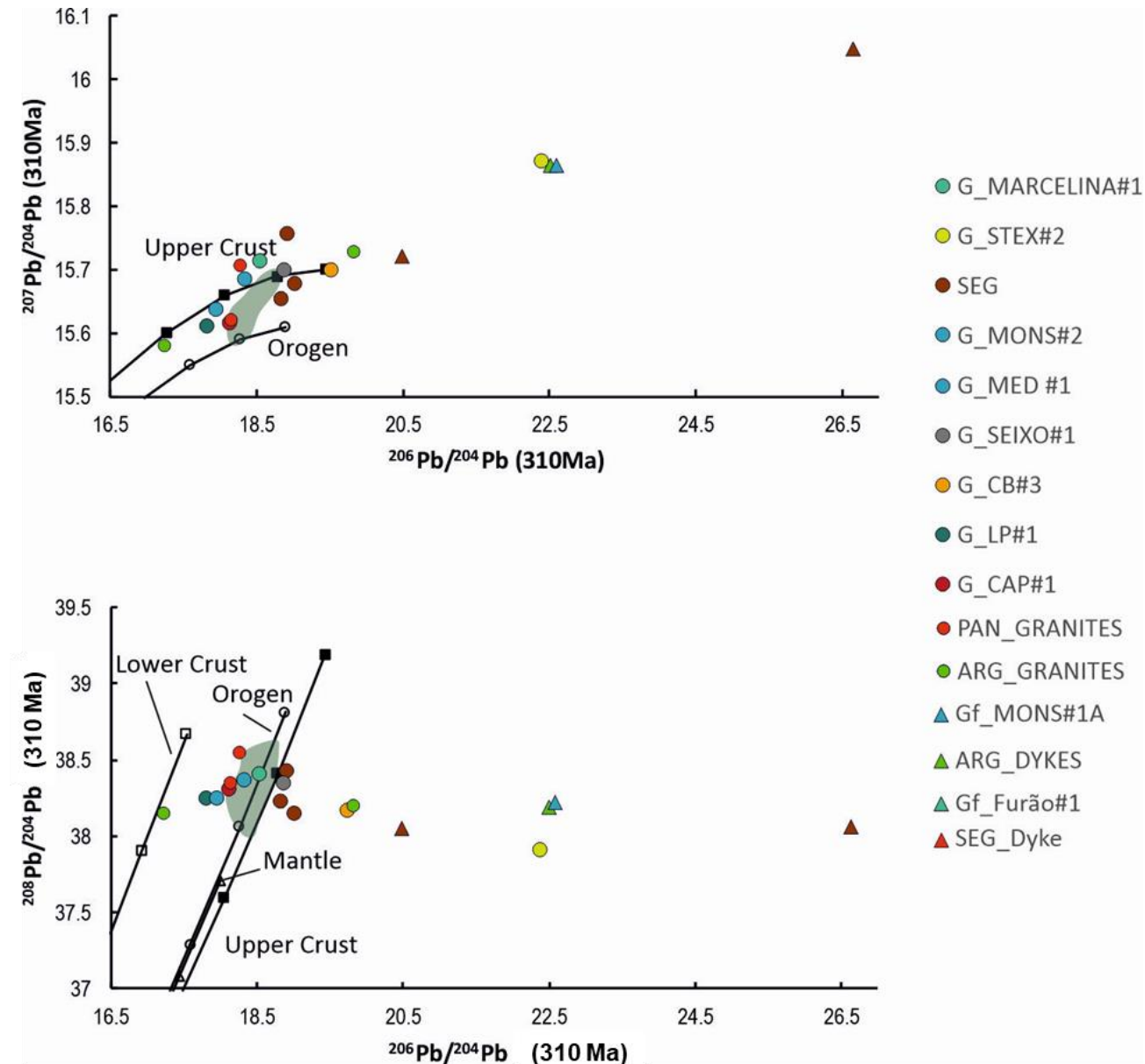




# Compositional features | Variscan magmatism

## Pb systematics

- **Pb-Pb isotopic ratios** spread between the “*orogen*” and “*upper crustal*” reference curves of the plumbotectonic model, but also plot above the “*upper crustal*” growth curve, displaying evident  $(^{207}\text{Pb}/^{204}\text{Pb})_{310\text{Ma}}$  and  $(^{206}\text{Pb}/^{204}\text{Pb})_{310\text{Ma}}$  increase.
- **The overlap with the Pb-Pb isotopic signature of the Beiras Group metapelites is limited**, indicating the involvement of other Pb sources, such as immature metasediments and/or meta-igneous protoliths.
- **The “Pb signature” for the Variscan granite suites supports a dominant upper crustal Pb derivation without significant juvenile contamination.**



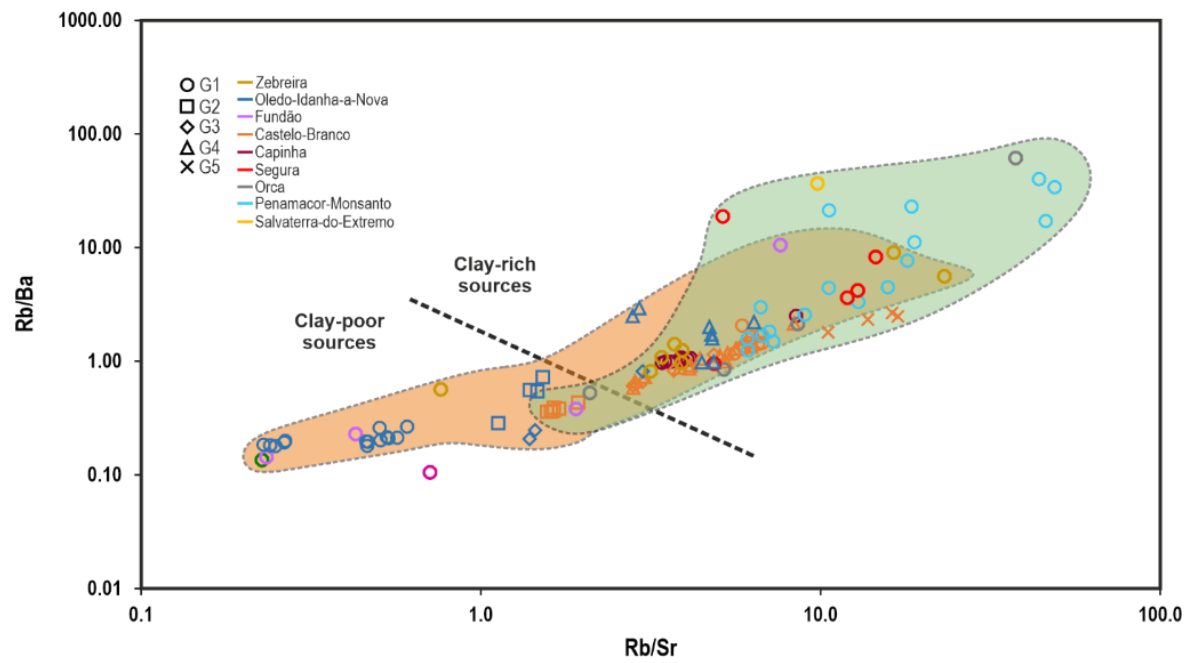
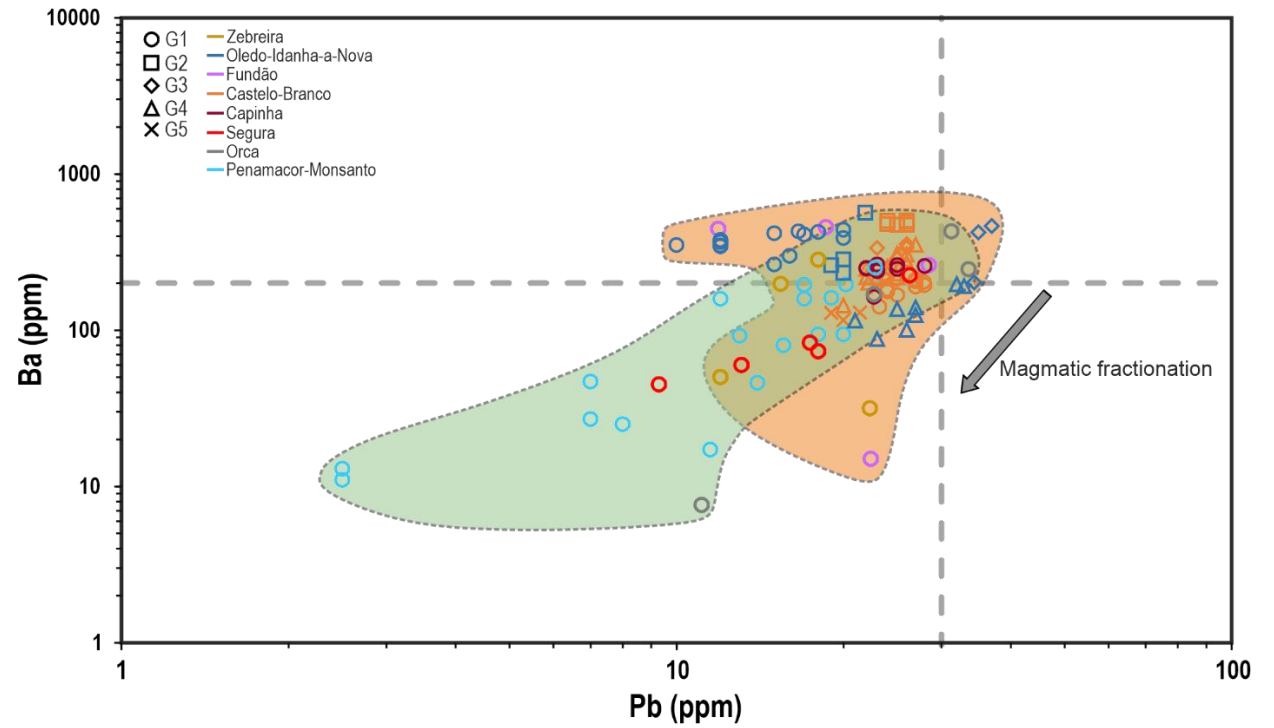
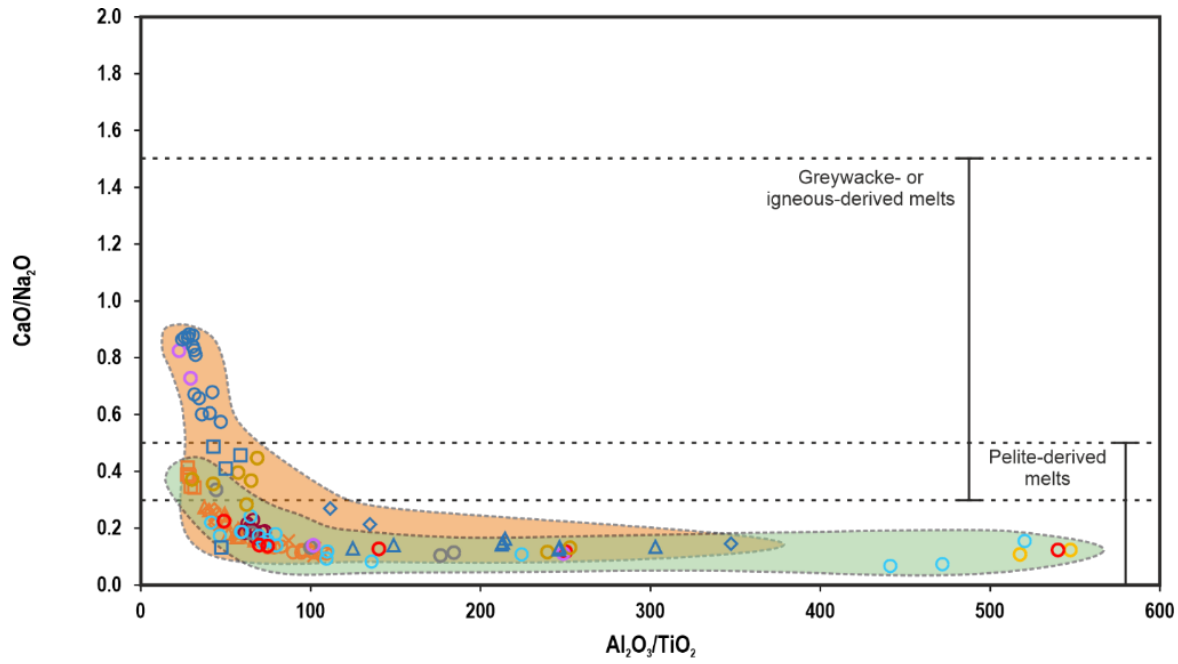


## Variscan magmatism

*(Panasqueira, Argemela, Castelo Branco, Orca, Capinha, Monsanto, Salvaterra do Extremo, Segura)*

- Multi-stage partial melting of different crustal levels.
- Some of these crustal levels also acted as protoliths of the previous Cambrian-Ordovician magmatic event, as suggested by  $\epsilon\text{Nd}$  values, Nd TDM model ages and Pb-Pb isotope ratios.





Finger and Schiller (2012):

- primary low-T granites;
- secondary low-T granites, fractionated from high-T parental magma (Pb < 30 ppm, Ba < 200 ppm).

Constraints to the estimation of zircon saturation temperatures ( $T_{\text{sat}}^{\text{zir}}$ ; Watson and Harrison, 1983)



**T<sub>zir</sub>  
sat (°C)**

**Temperature conditions for the extraction of a granitic melt from its source assuming no significant fractional crystallisation has occurred**

**Cambrian-Ordovician granites**

Facies	Zebreira		OIN-G1		OIN-G2	OIN-G3		Fundão	
Sample	G_ZEB#1	G_ZEB#2	GPF	OLX	G_IDN#1	REPA1	REPA2	G_FUN#2	G_FUN#3
T (°C)	761	765	754	764	755	718	720	742	737

**Variscan granites**

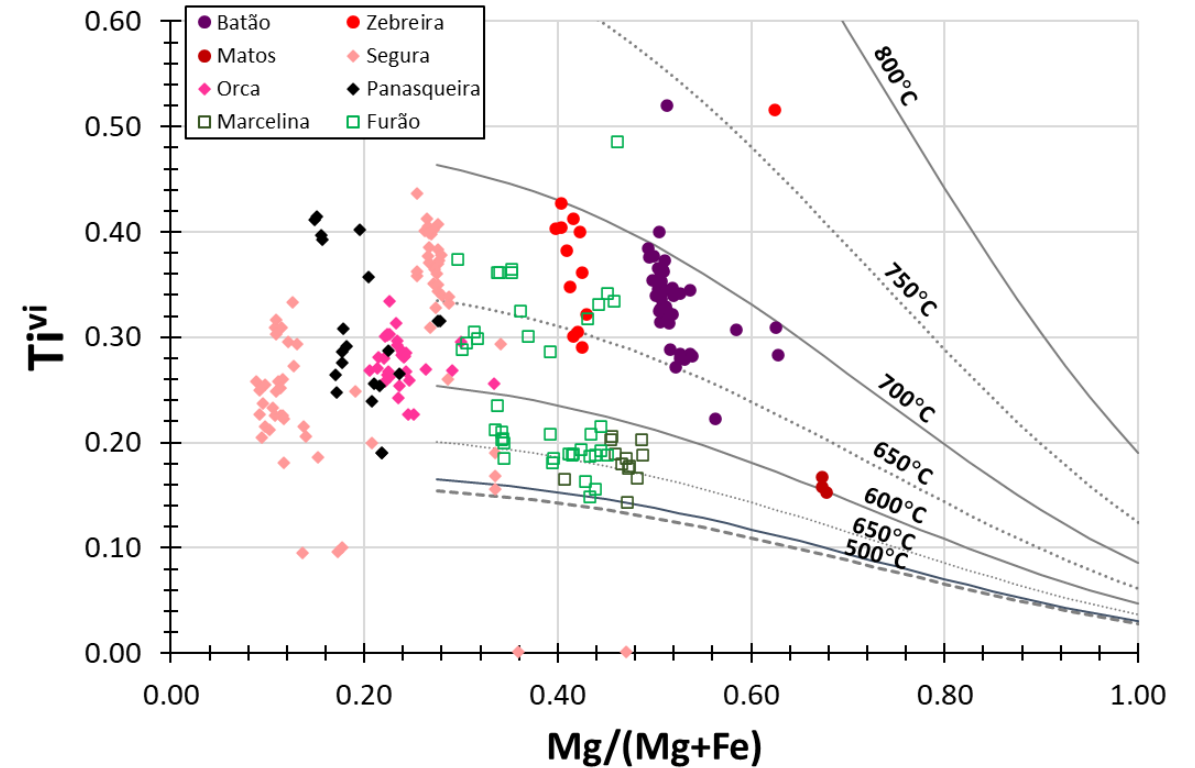
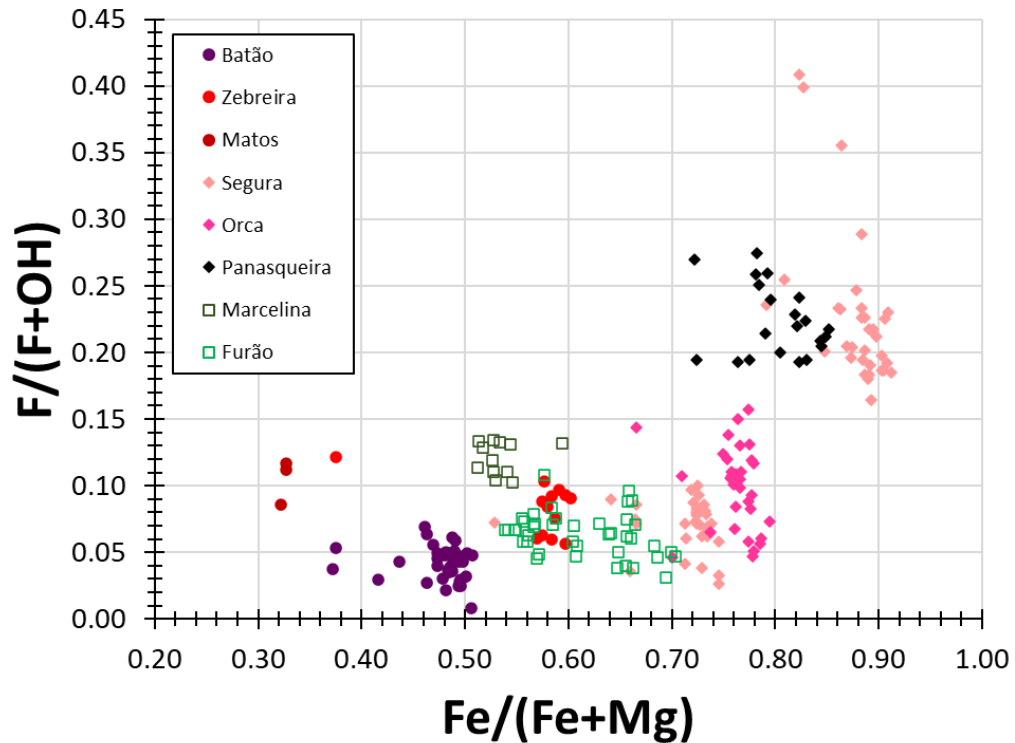
Facies	CB-G2						SEG	P-M		Orca	
Sample	GIN	GIN2	GIN4	G14	INFX1	INFX2	G_SEG#4	G_MONS#2	P-39	G_STC#1	G_MDCH#1
T (°C)	824	820	821	818	821	813	767	763	792	759	712

**Capinha**

Facies	Capinha										
Sample	G_CAP#1	1	4	11	15	18	19	Q1	Q2		
T (°C)	750	753	755	749	757	744	749	752	751		

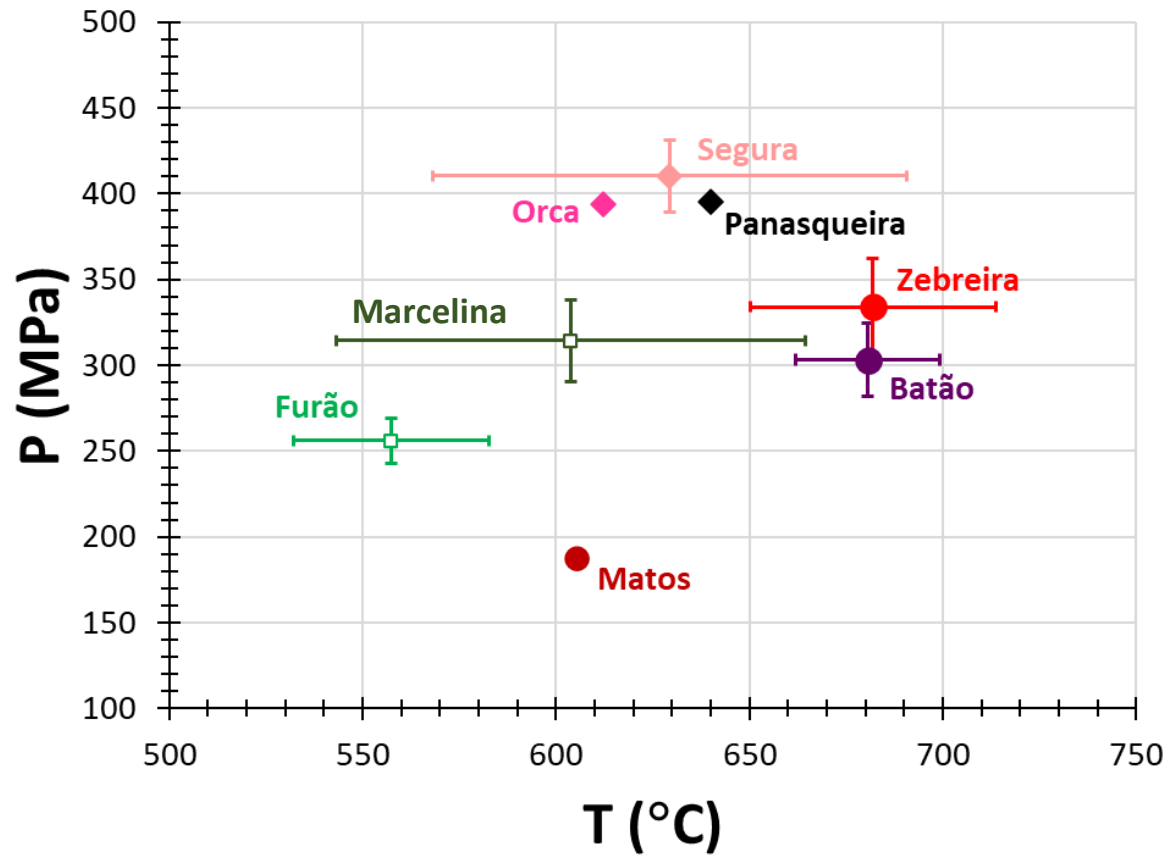
# Biotite | ongoing work

- Protracted crystallisation path
- Usual interstitial growth (in relation to alkali feldspar, quartz, and plagioclase), suggesting similar timing of biotite saturation
- Biotite composition should record conditions at comparable extents of magma evolution for different rock suites.

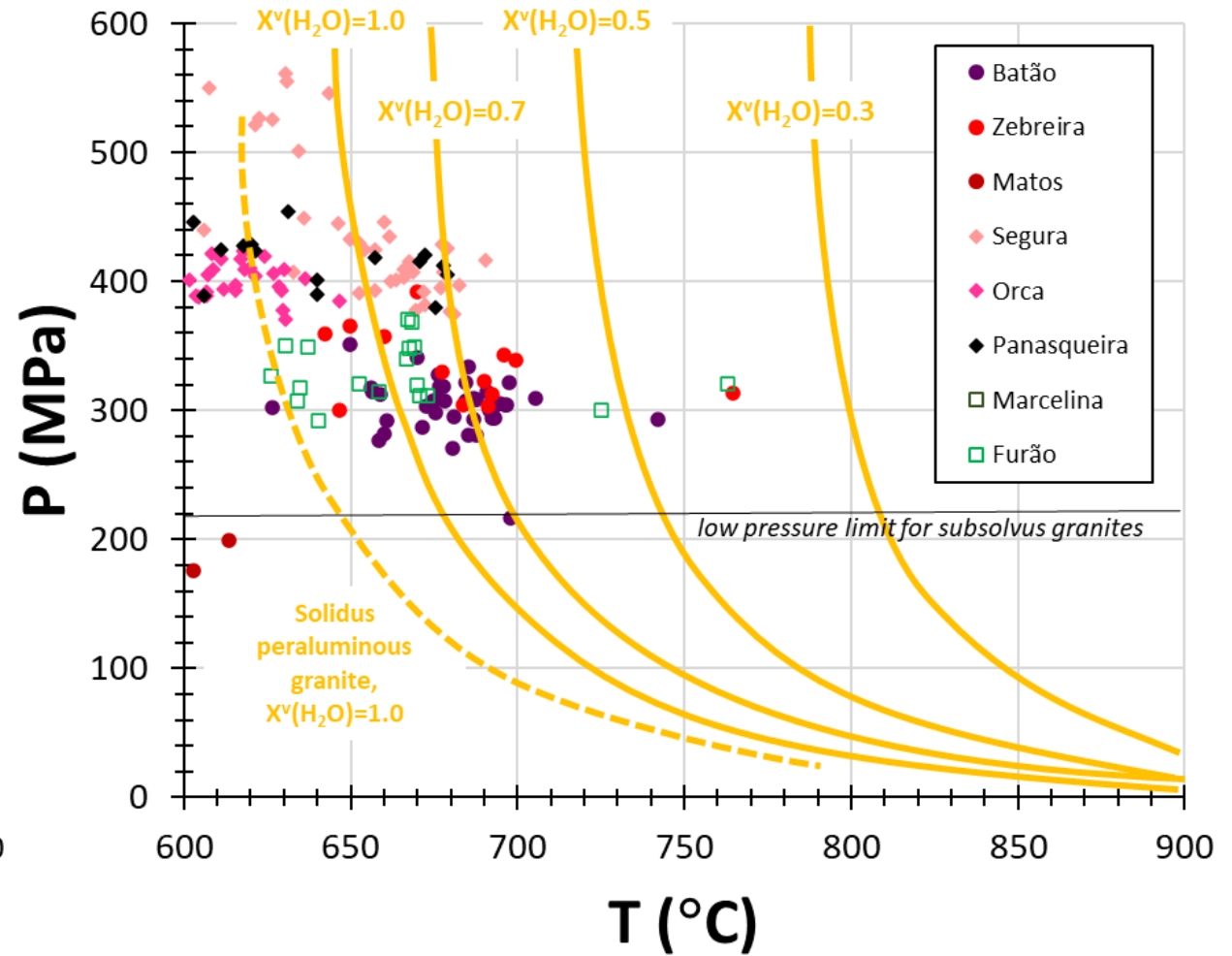


- Biotite T-crystallisation based on Ti contents and  $Mg/(Mg+Fe)$  (Henry et al., 2005)
- Application limited to  $0.04 \leq Ti \leq 0.60$  apfu and  $0.275 \leq Mg/(Mg+Fe) \leq 0.60$

# Biotite | ongoing work



P estimates using the Al content and the relationship of Henry et al. (2005)



Minimum water contents above 6-7 wt% for the magmas, conceivably higher in Variscan suites in comparison to those generated during the Cambrian-Ordovician transition



# Geochronology | starting point

Antunes et al. (2009)

- Oledo – Idanha-a-Nova**
- “Granodiorite G1”: ID-TIMS U-Pb zircon: 480.5±1.0 Ma; ID-TIMS U-Pb monazite: 478.3±1.1 Ma
  - “Granodiorite G2”:
  - “Granodiorite G3”: ID-TIMS U-Pb zircon: 479.0±4.0 Ma
  - “G4 γ”: ID-TIMS U-Pb zircon: 479.0±3 Ma

Several studies; see Mateus et al. (2020)

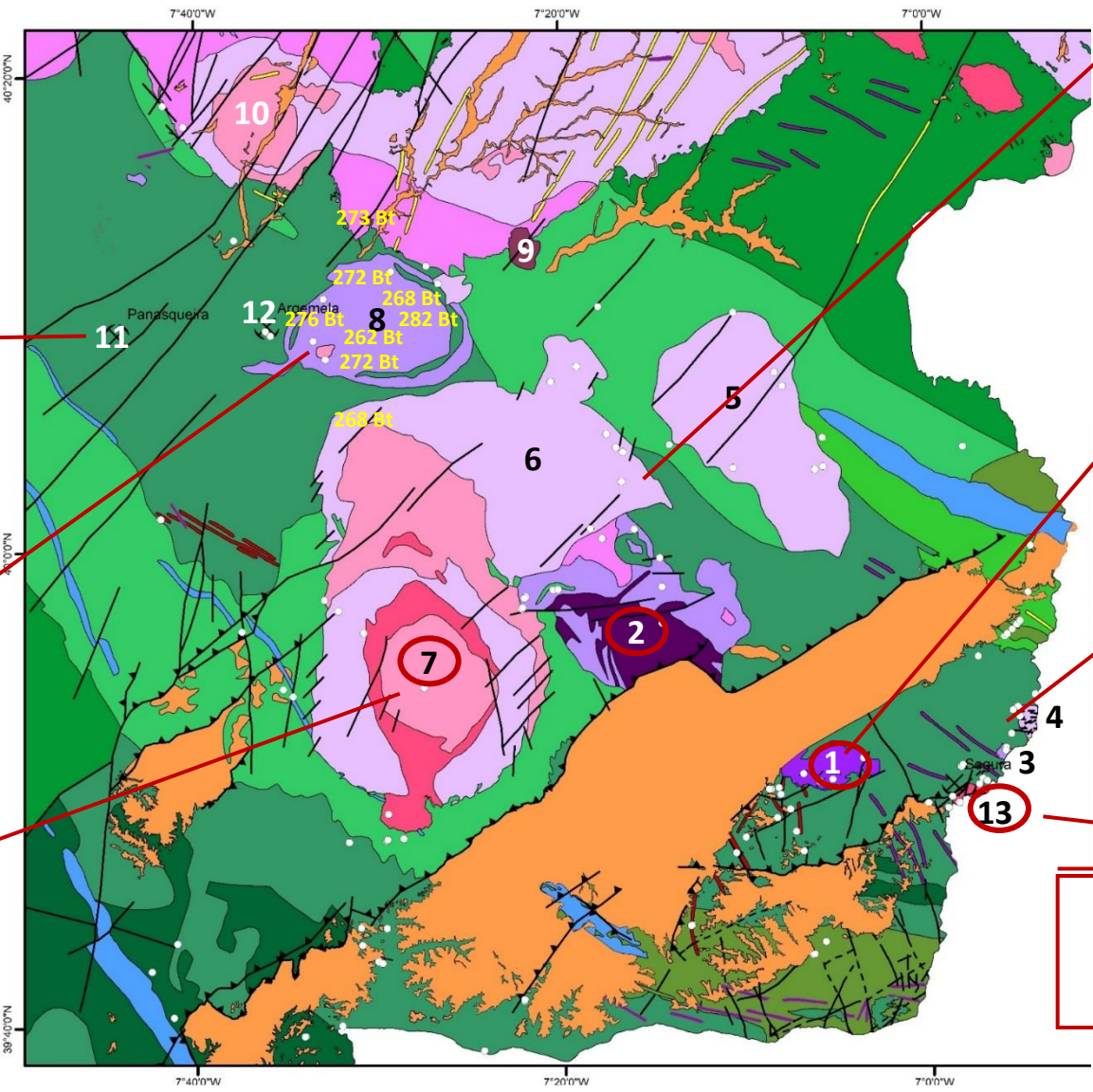
- Panasqueira**
- WR Rb-Sr, two-mica γ: 289±4 Ma
  - U-Pb uraninite in greisen: 298±2 Ma
  - U-Pb rutile (border of qz-lodes): 305±3.3 Ma
  - U-Pb cassiterite (qz-lodes): 303±3.3 Ma and 301±4.2 Ma
  - Ar-Ar muscovite: 296.3±0.8 Ma (thick selvages), 291.6±0.8 Ma (relative late mica)

Portugal Ferreira et al. (1977)

- Fundão (Covilhã, Alpedrinha)**
- K-Ar biotite: 262±2 Ma e 282±2 Ma (different facies forming the Fundão pluton)
  - K-Ar biotite: ≈270 Ma porphyroid γs of Covilhã and Alpedrinha

Antunes et al. (2008)

- CASTELO BRANCO**
- “Ms>Bt γ core”: ID-TIMS U-Pb zircon: 309.0±1.1 Ma; ID-TIMS U-Pb monazite: 309.5±0.9 Ma
  - “Bt>Ms granodiorite ring”: ID-TIMS U-Pb zircon: 310.1±0.8 Ma; ID-TIMS U-Pb monazite: 310.5±1.5 Ma
  - Porphyritic “Bt>Ms granodiorite” grading to Bt≡Ms γ
  - “Ms>Bt external γ ring”: ID-TIMS U-Pb zircon: 309.7±0.4 Ma; ID-TIMS U-Pb monazite: 309.0±3.2 Ma



Pereira et al. (1986)

- Zebreira**
- “Granodiorite”: K-Ar biotite: 372±7 Ma; K-Ar muscovite: 472±7 Ma
  - “Monzo-γ dykes”: K-Ar muscovite: 440±8 Ma

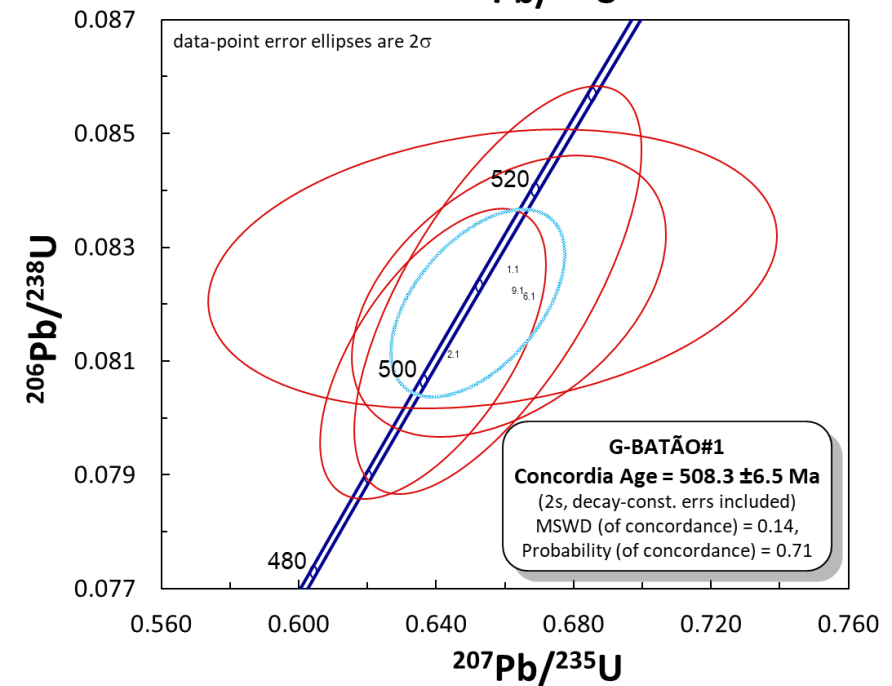
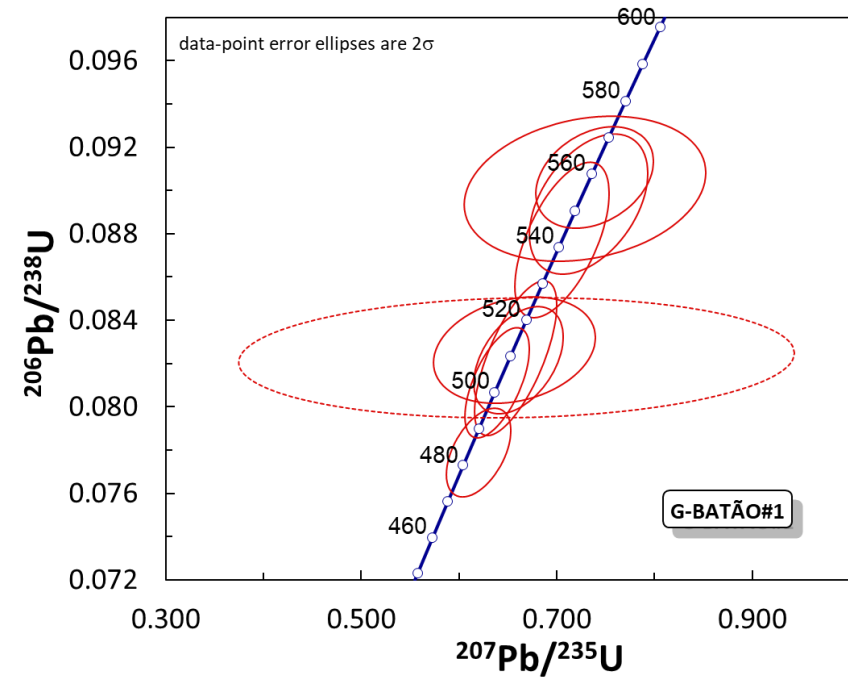
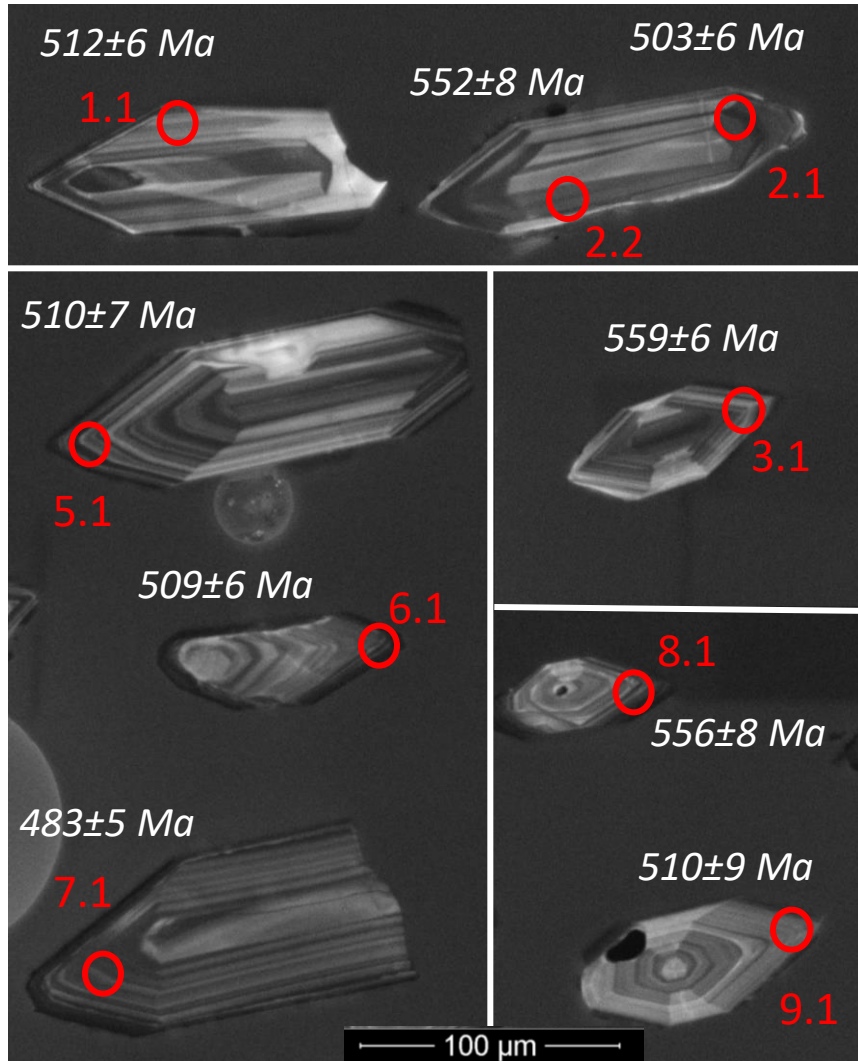
Ribeiro et al. (2010)

- Salvaterra do Extremo area**
- Foliated mafic dykes, ascribed to the Cambrian-Ordovician magmatic event

Antunes et al. (2013)

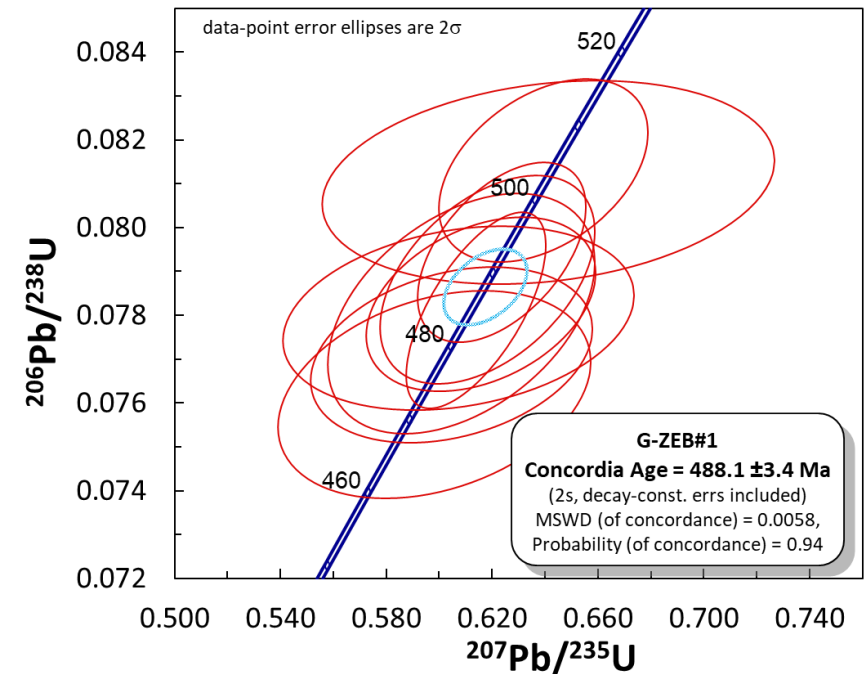
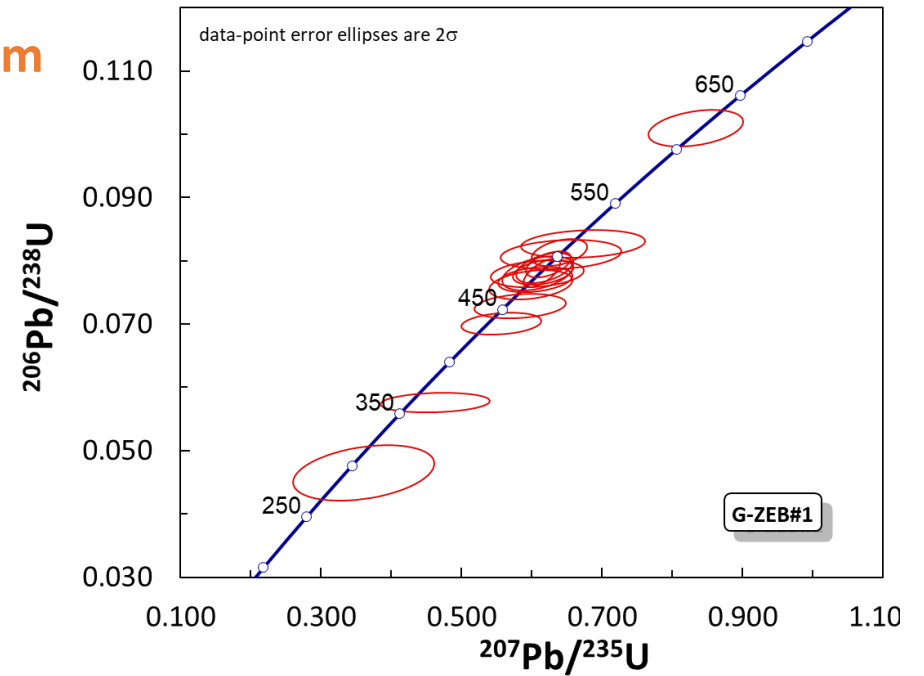
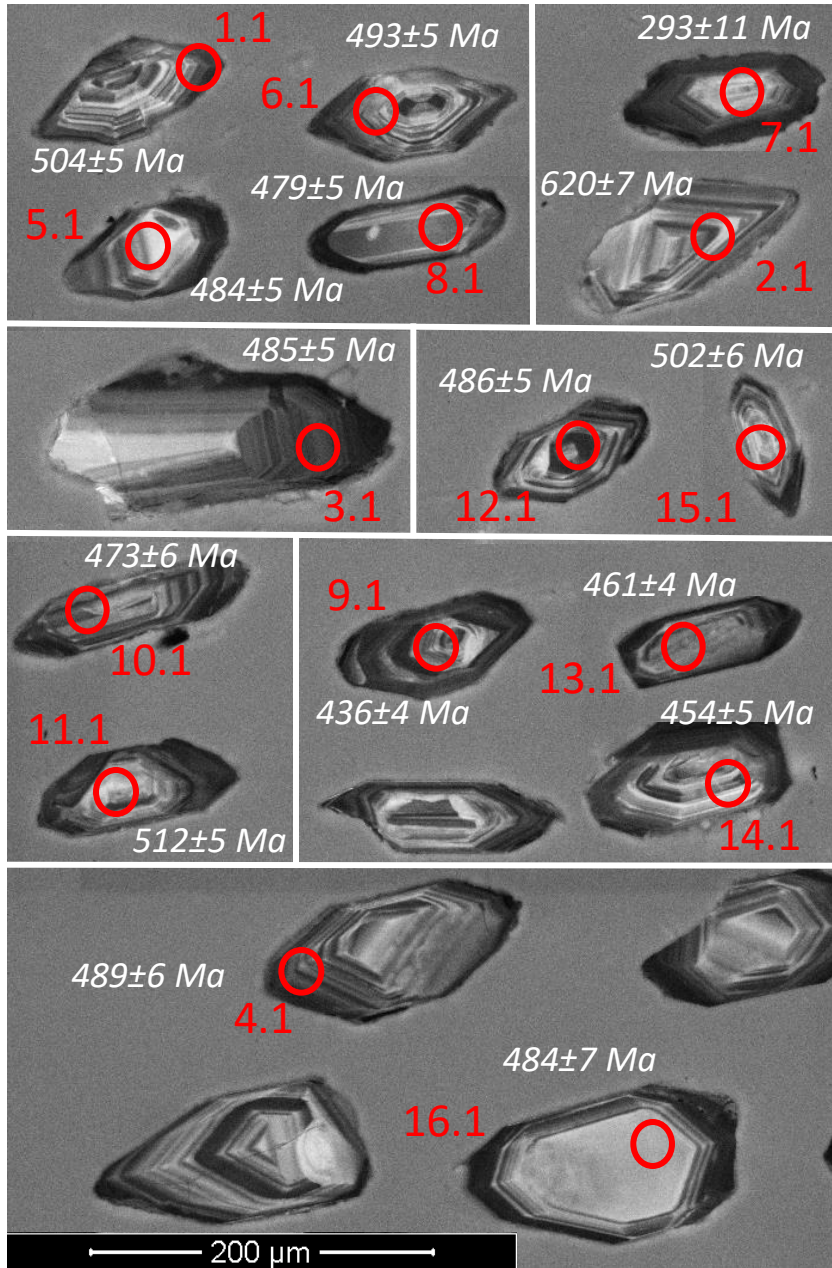
- SEGURA**
- “Two-mica γ”: ID-TIMS U-Pb zircon: 311.0±0.5 Ma; ID-TIMS U-Pb monazite: 312.9±2.3 Ma
  - “Ms-γ”: ID-TIMS U-Pb zircon: 312.9±2.0 Ma

# Batão tonalite | Cambrian-Ordovician magmatism

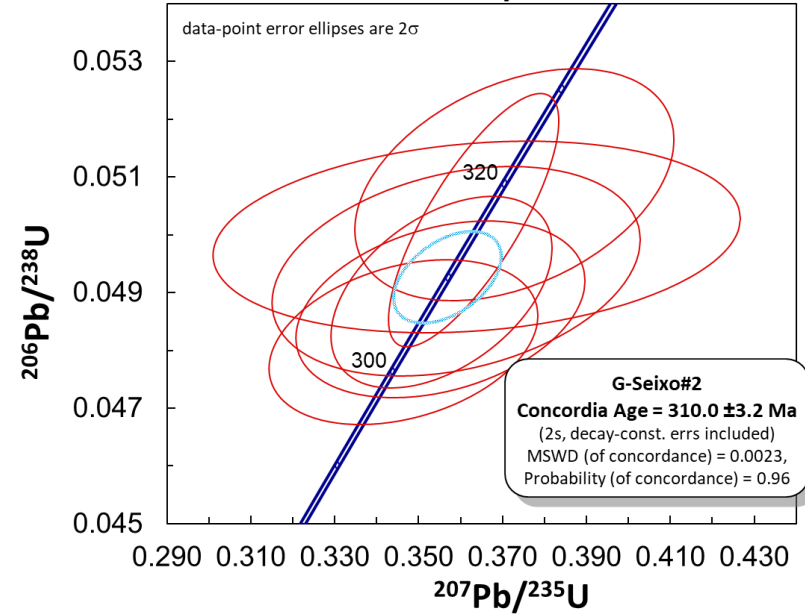
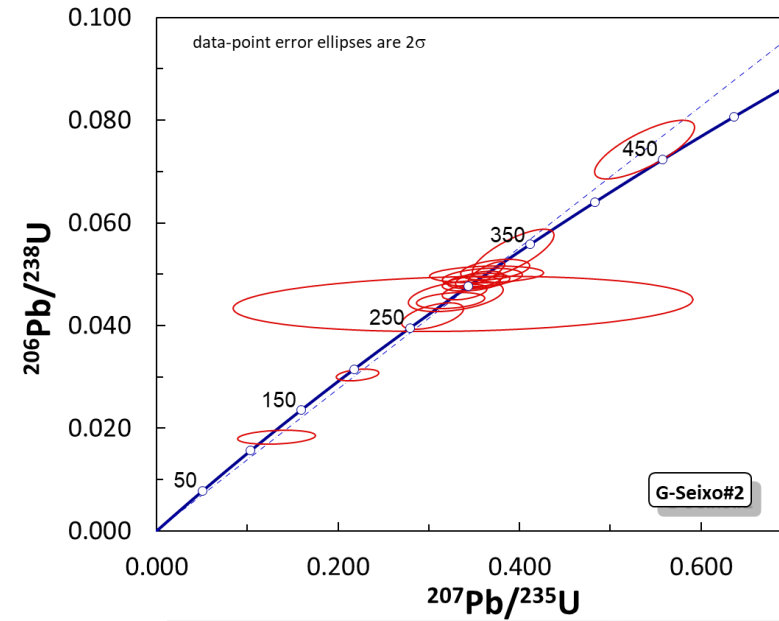
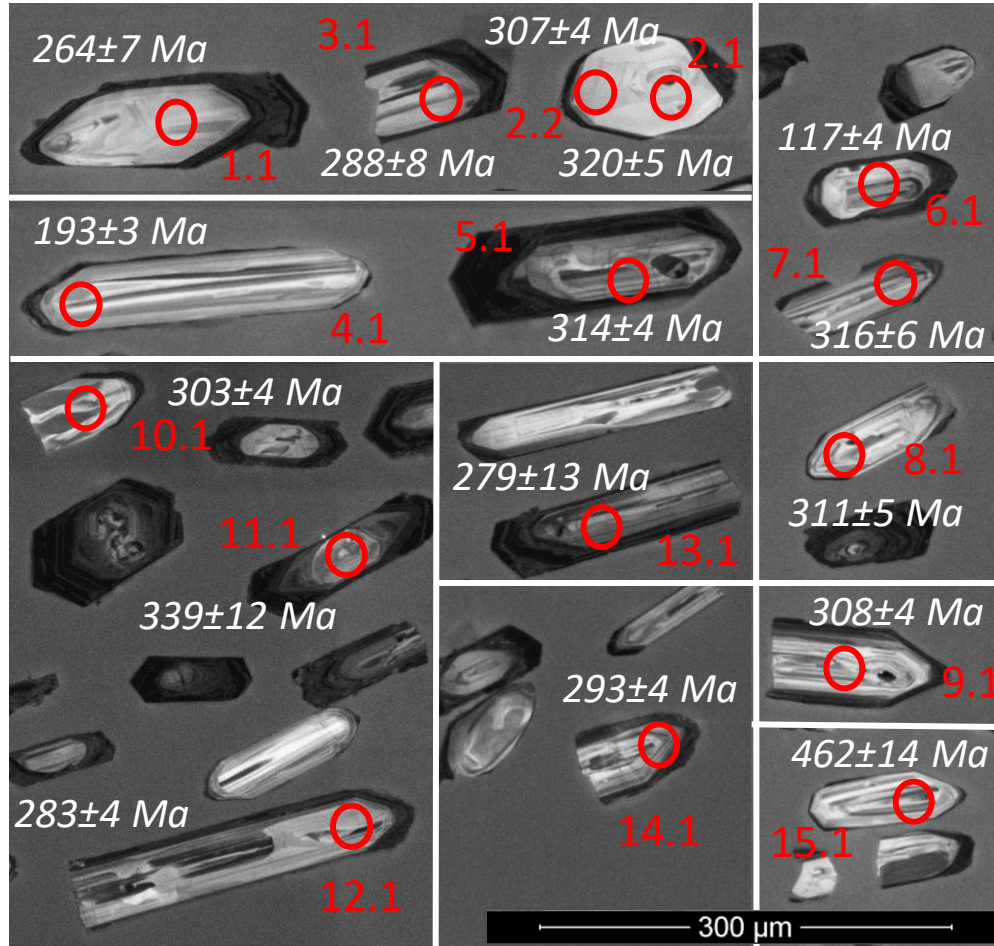




# Zebreira granodiorite | Cambrian-Ordovician magmatism

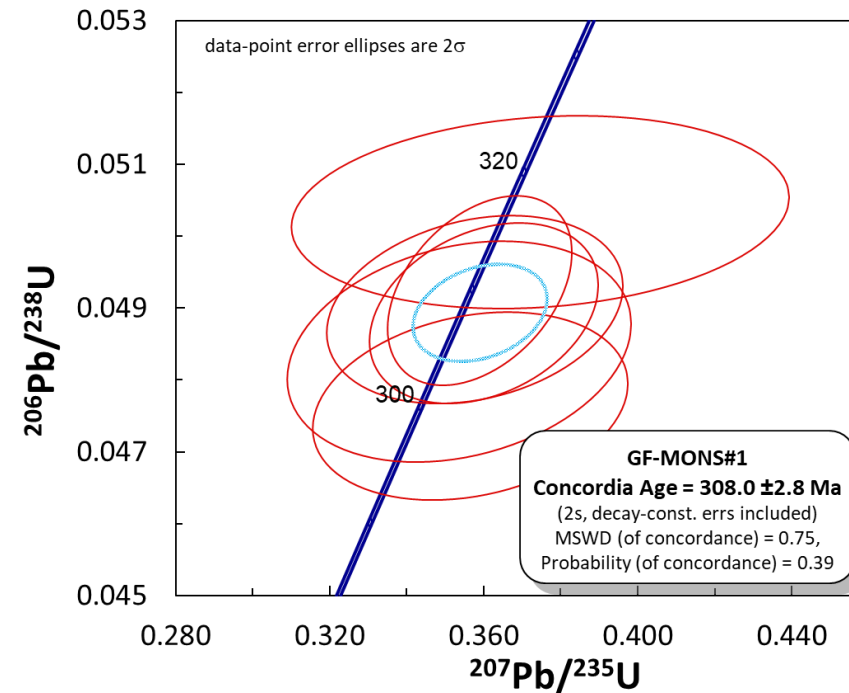
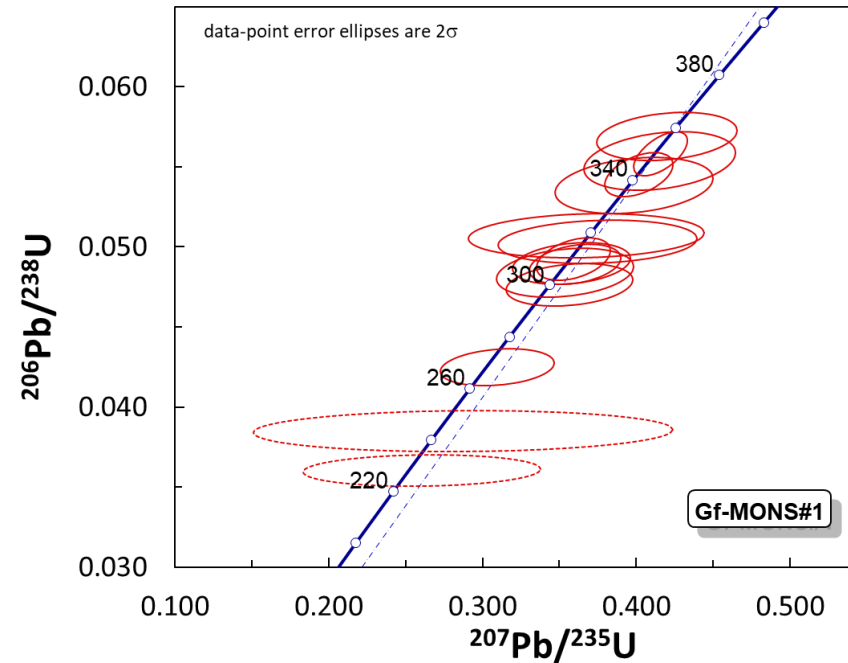
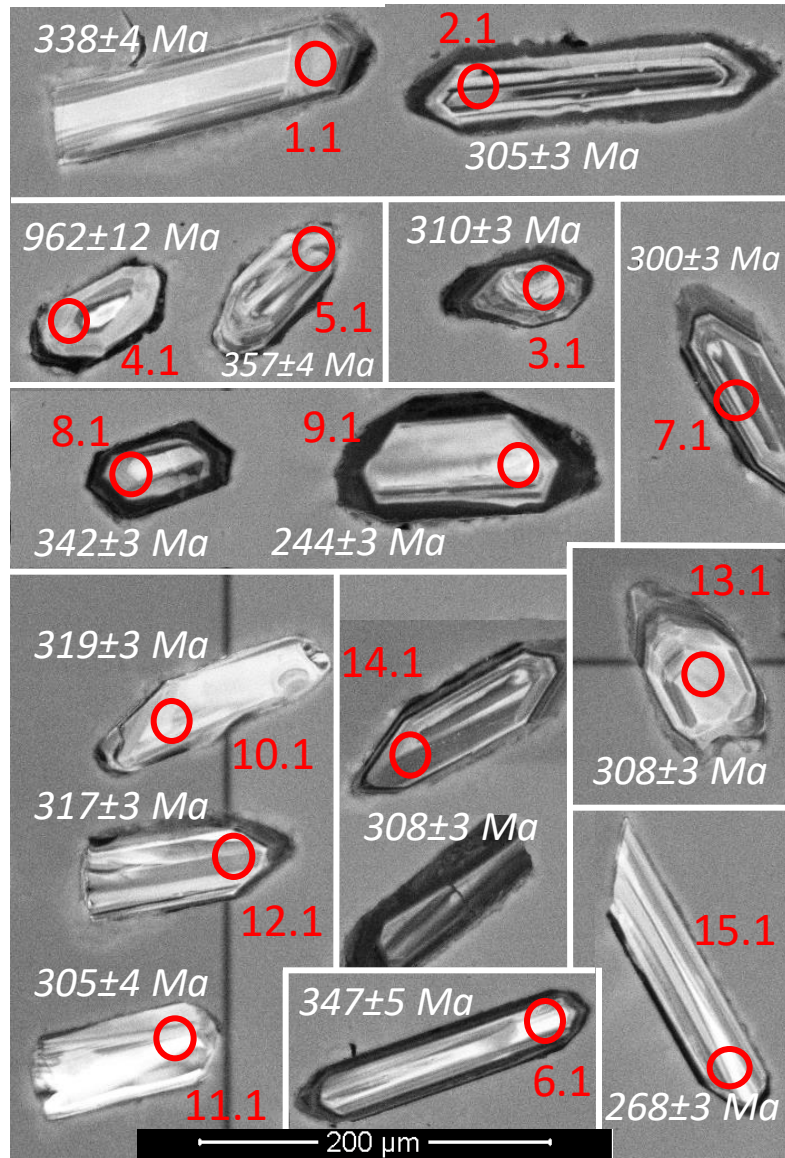


# Orca granite | Variscan magmatism

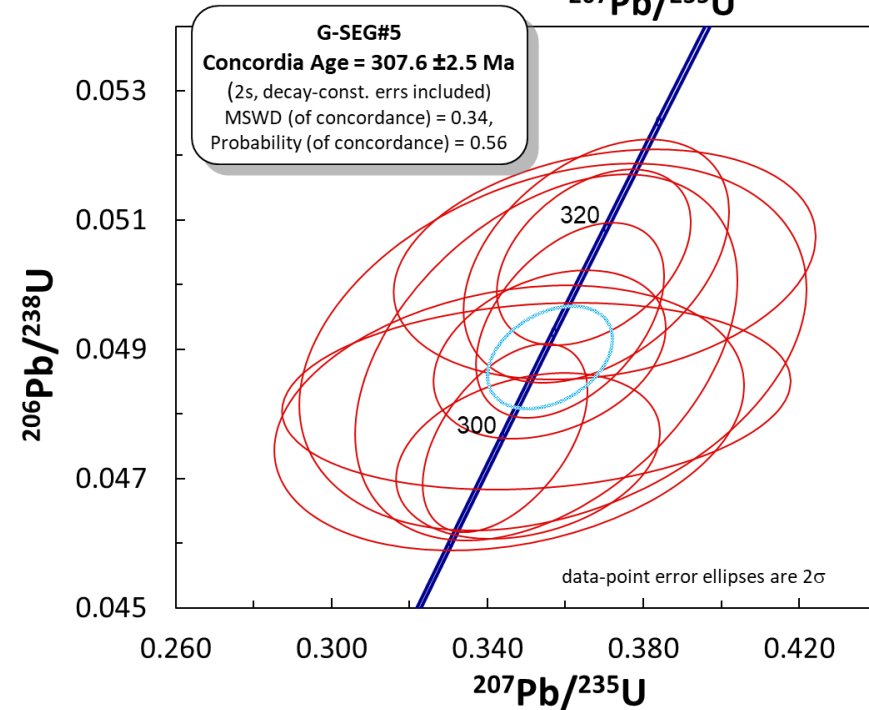
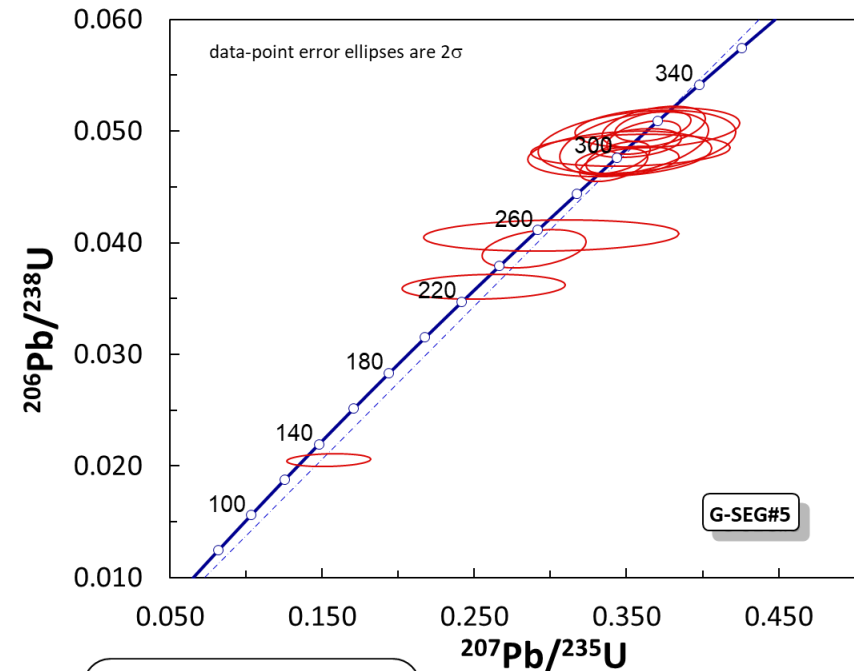
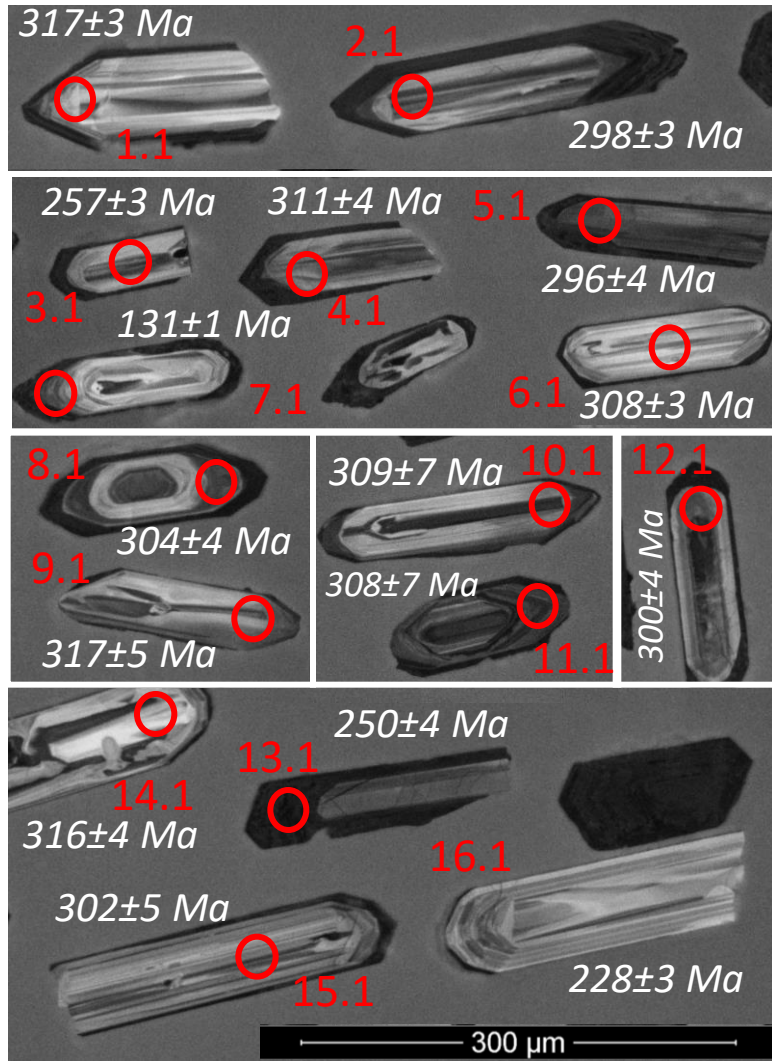




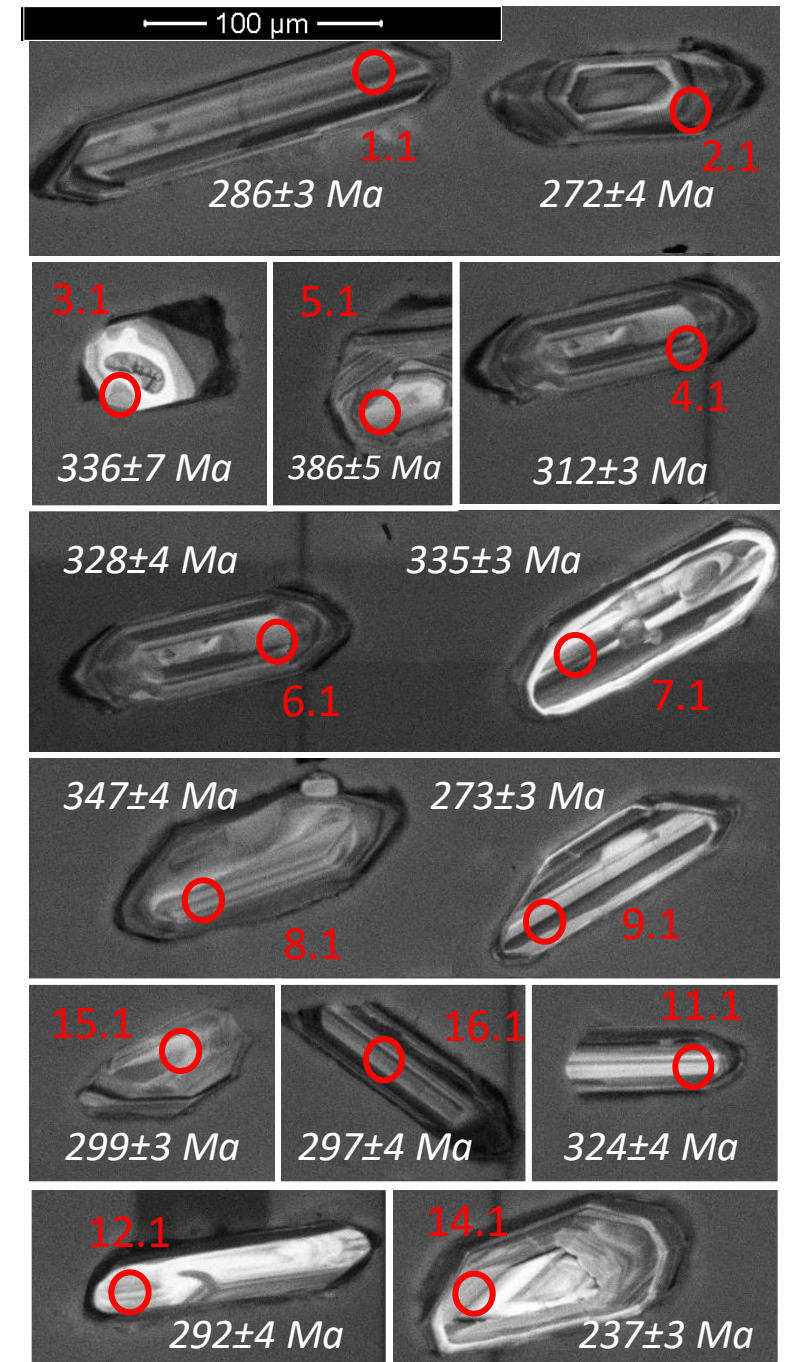
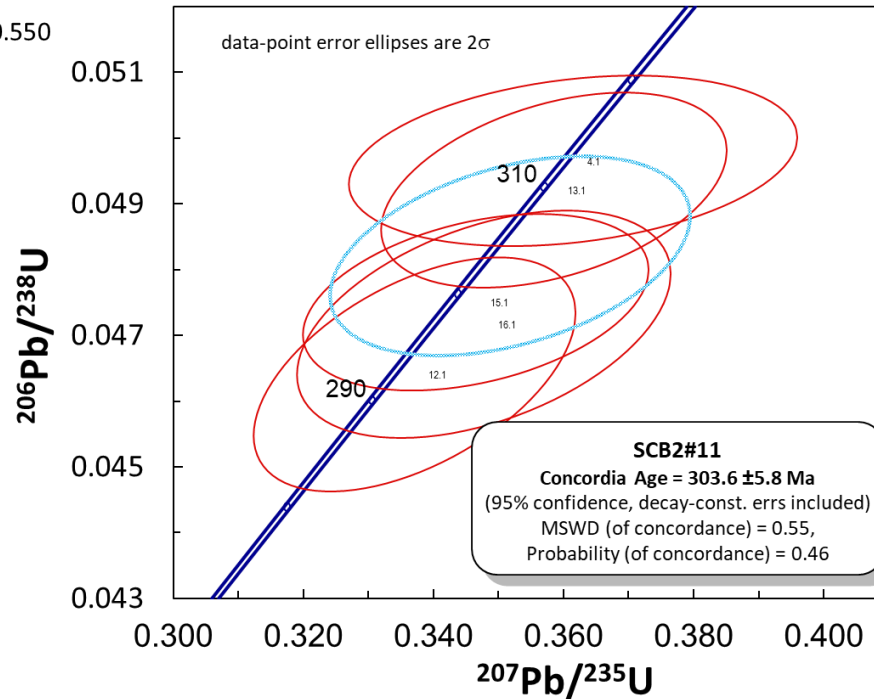
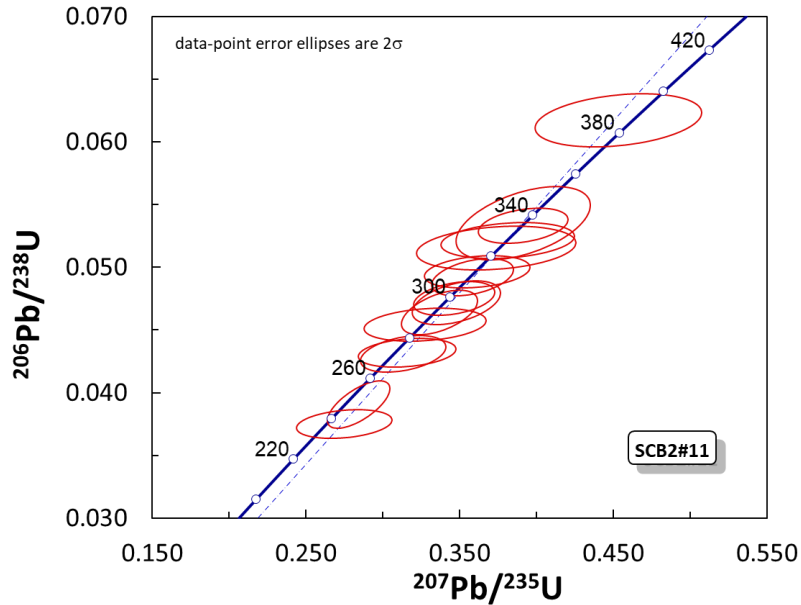
# Monsanto pegmatite | Variscan magmatism



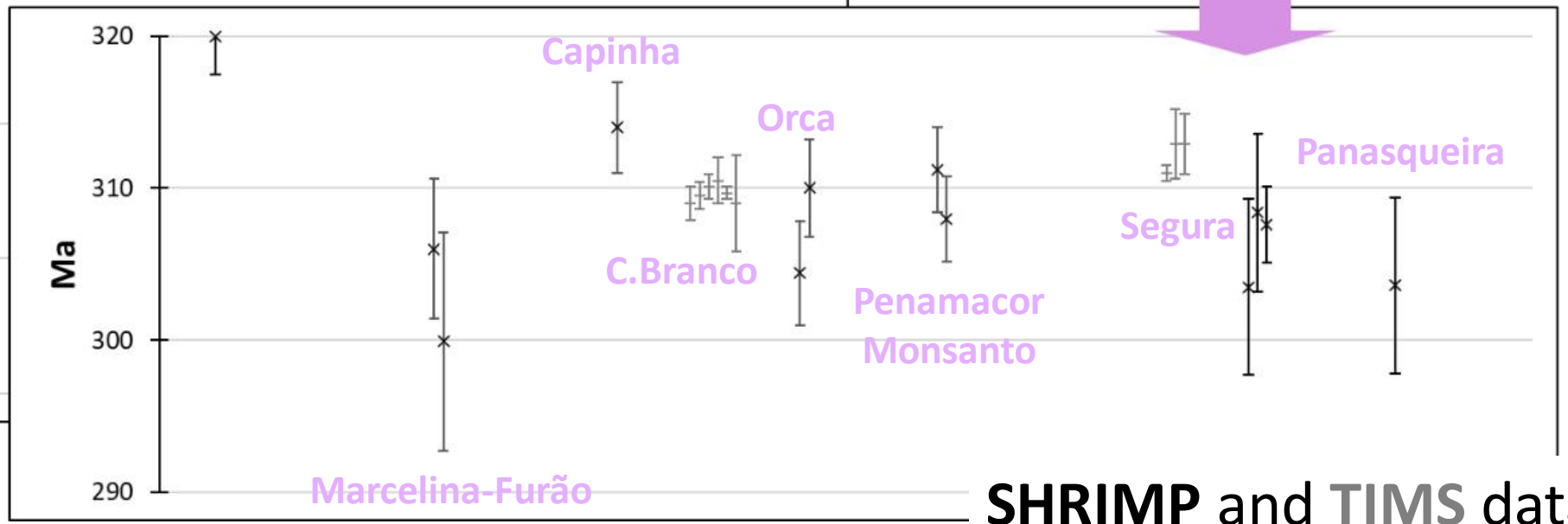
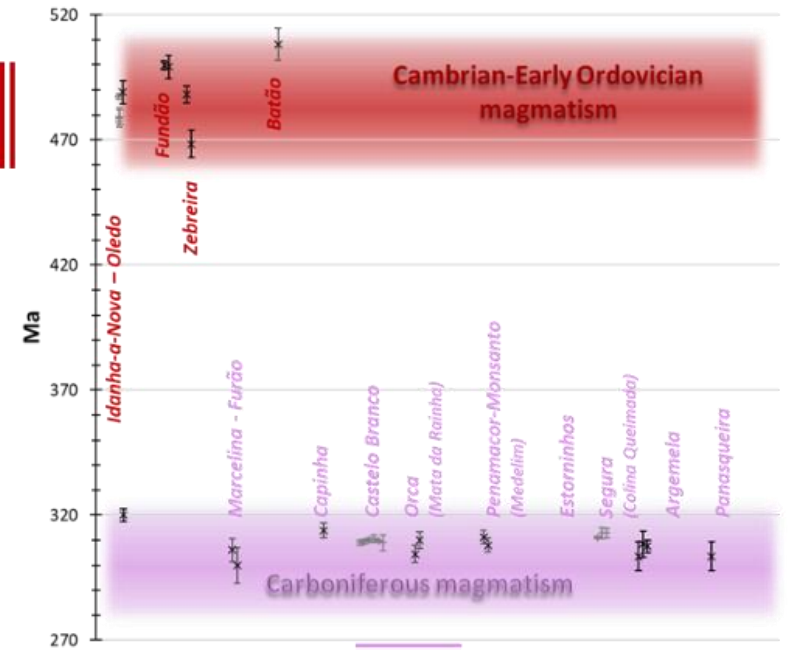
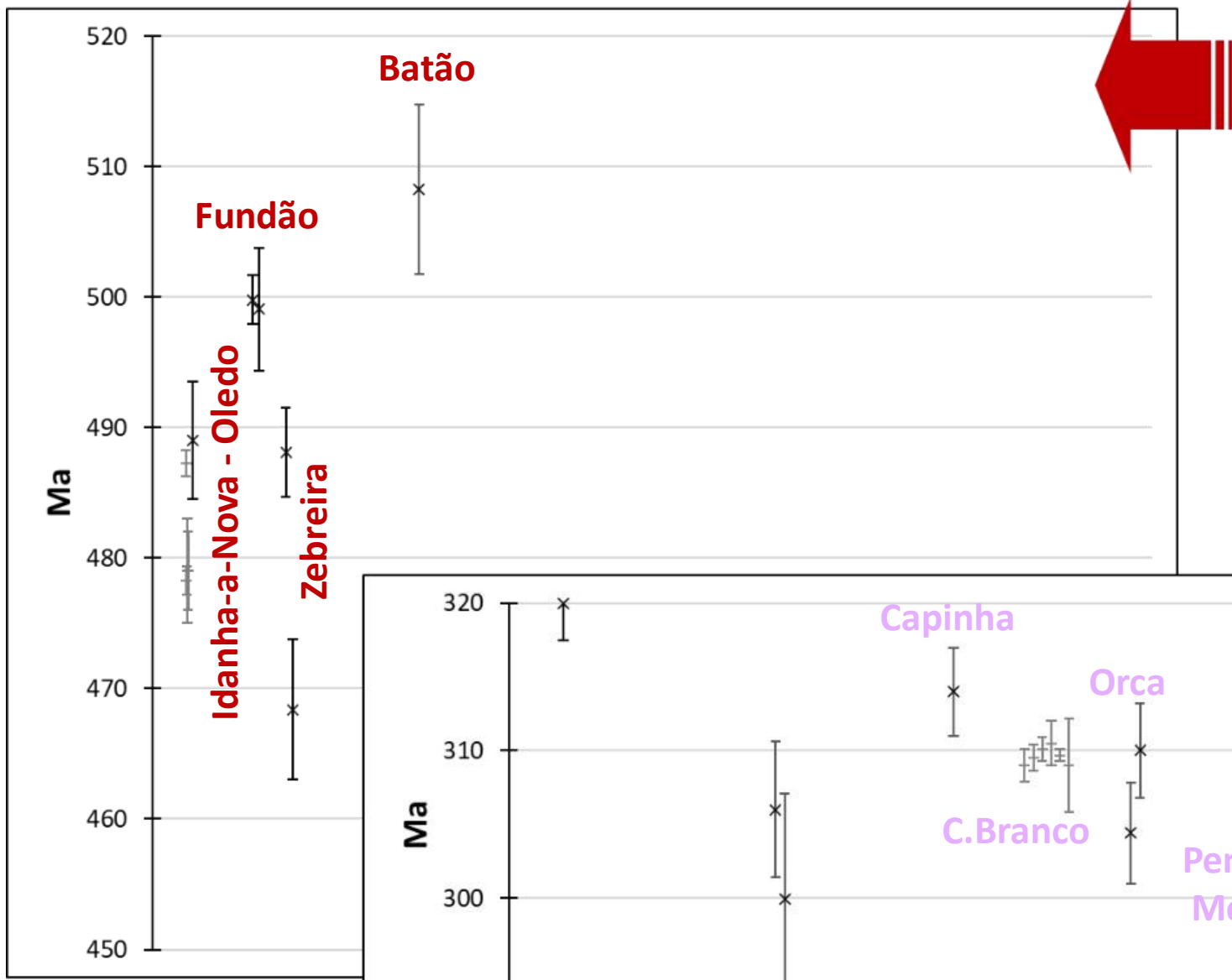
# Segura *ms*-granite | Variscan magmatism



# Panasqueira two-mica granite | Variscan magmatism







**SHRIMP and TIMS data**



# Geochronology | *SHRIMP U-Pb zircon*

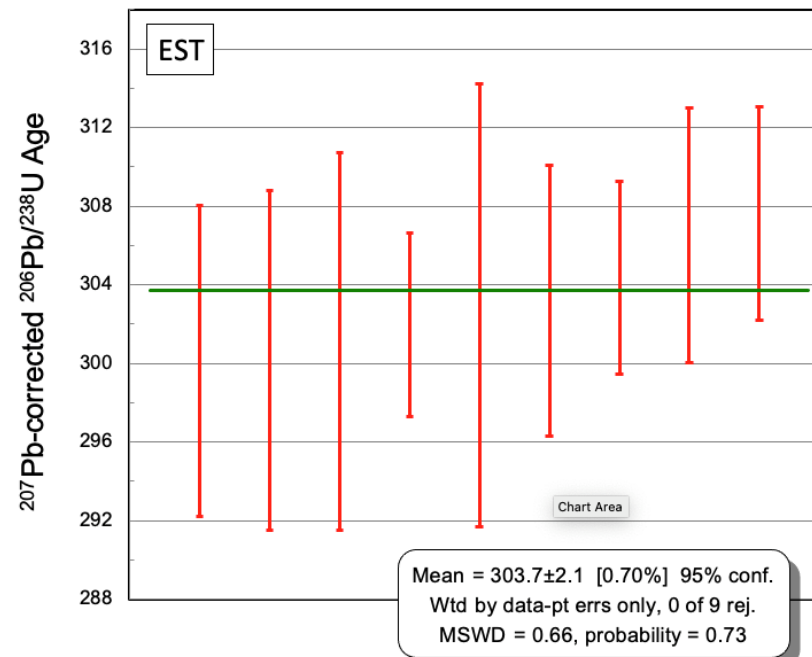
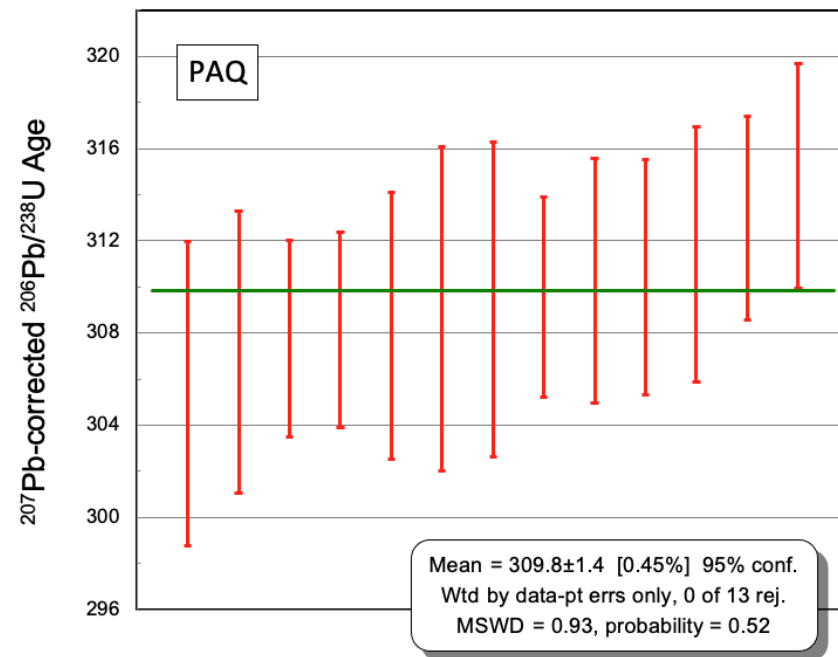
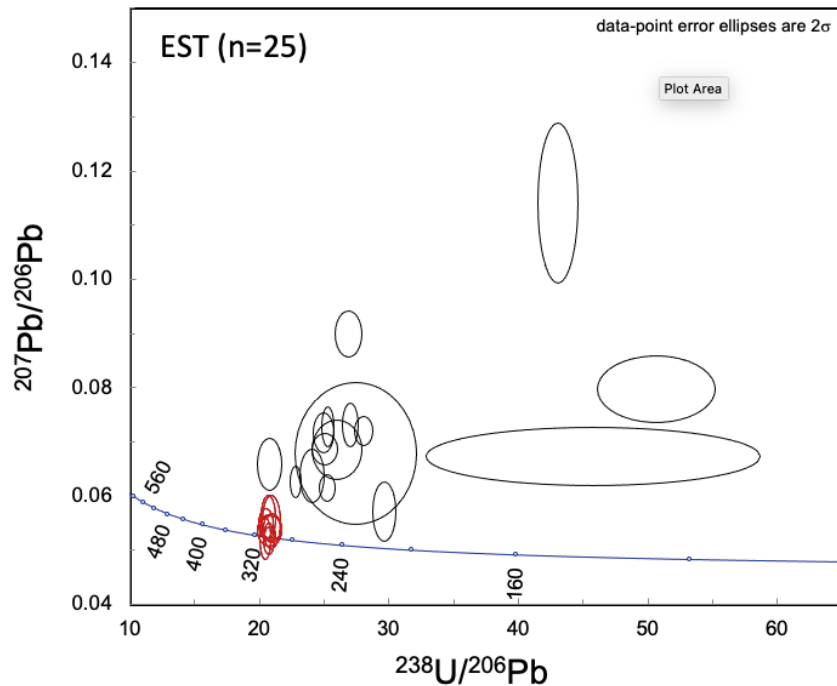
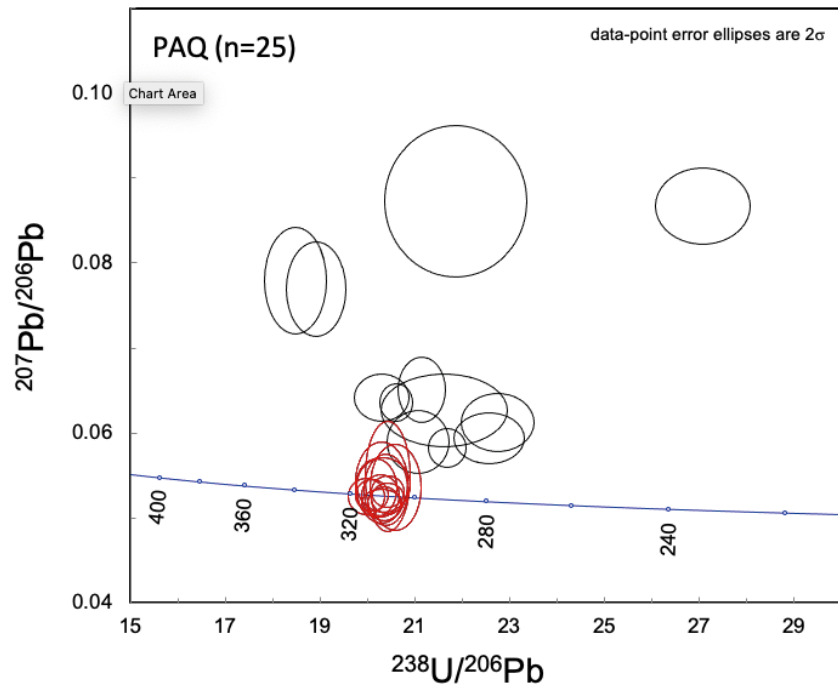
Notwithstanding the good match of results, several problems have come up with different samples selected for geochronological studies, due to:

- Extreme difficulty, or even impossibility, to obtain zircon concentrates in various highly differentiated or compositionally modified granite facies from, e.g., Panasqueira, Argemela, and Medelim;
- Abundant inherited zircon populations in many granite facies, hindering a robust dating and often yielding ages with poor representativity/statistics of data;
- Insufficient zircon overgrowth encircling inherited cores or late rims with adequate thickness but containing too much uranium;
- Pb and even U loss, generating abundant discordant data, most of the times in zircon populations that include grains with evident effects of metamict processes; and
- The use of TEMORA standard in SHRIMP measurements (natural zircon with low to “normal” U content), possibly not so suitable for dating high-U zircons as those usually present in many of the sampled plutonic and dyke granite facies.

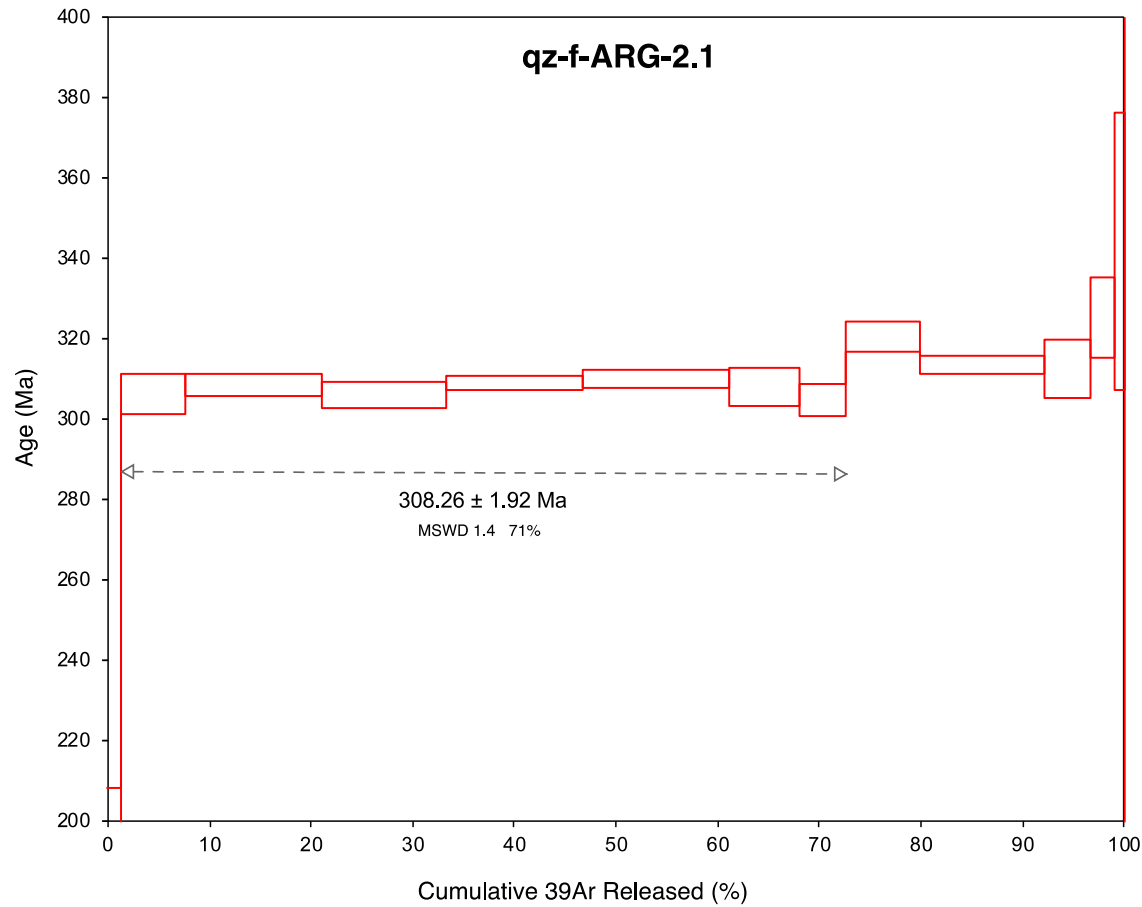
# Geochronology | *complementary approaches*

To overcome these difficulties and better constrain the temporal development of “fertile” granite melts and related mineralisation in the Góis-Panasqueira-Argemela-Segura strip, it was decided:

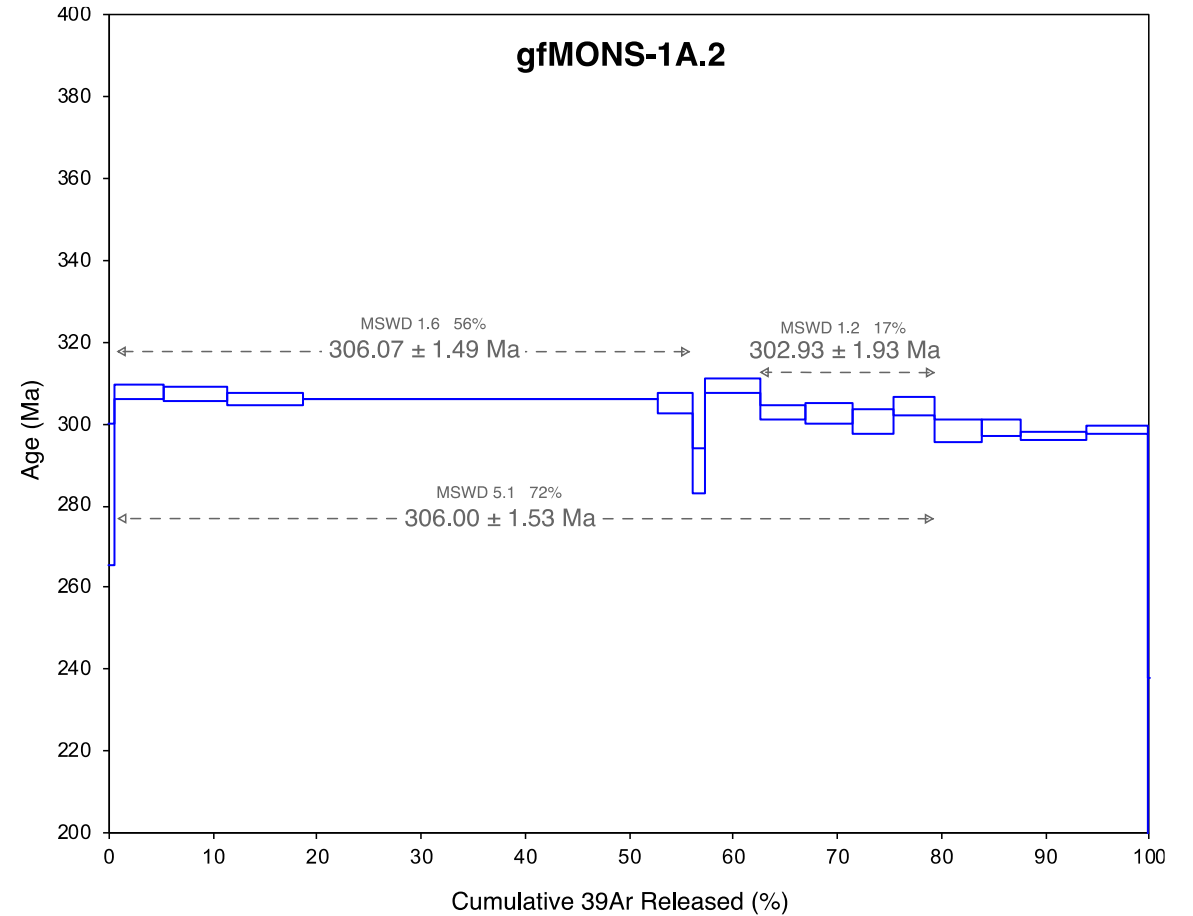
- to use LA-ICP-MS U-Pb zircon data to estimate the age of the cordierite facies of the Cabeza de Araya batholith (PAQ – Piedras Albas Quarry, the core domain of the Segura pluton) and the Estorniños (EST) two-mica granite; and
- to complement the available U-Pb zircon data with K-Ar and Ar-Ar ages of muscovite extracted from a 28 samples collected in different sites.
- The K-Ar ages are roughly equivalent to the whole age (WMA) using the Ar-Ar technique when plateau ages cannot be obtained on the whole spectra. The main advantage of dating muscovite grains with these two techniques results in the **estimation of the granite/aplite/pegmatite minimal age**, when the use of other approaches (such as U-Pb in zircon) is impossible or generates unlikely results from the geological point of view.







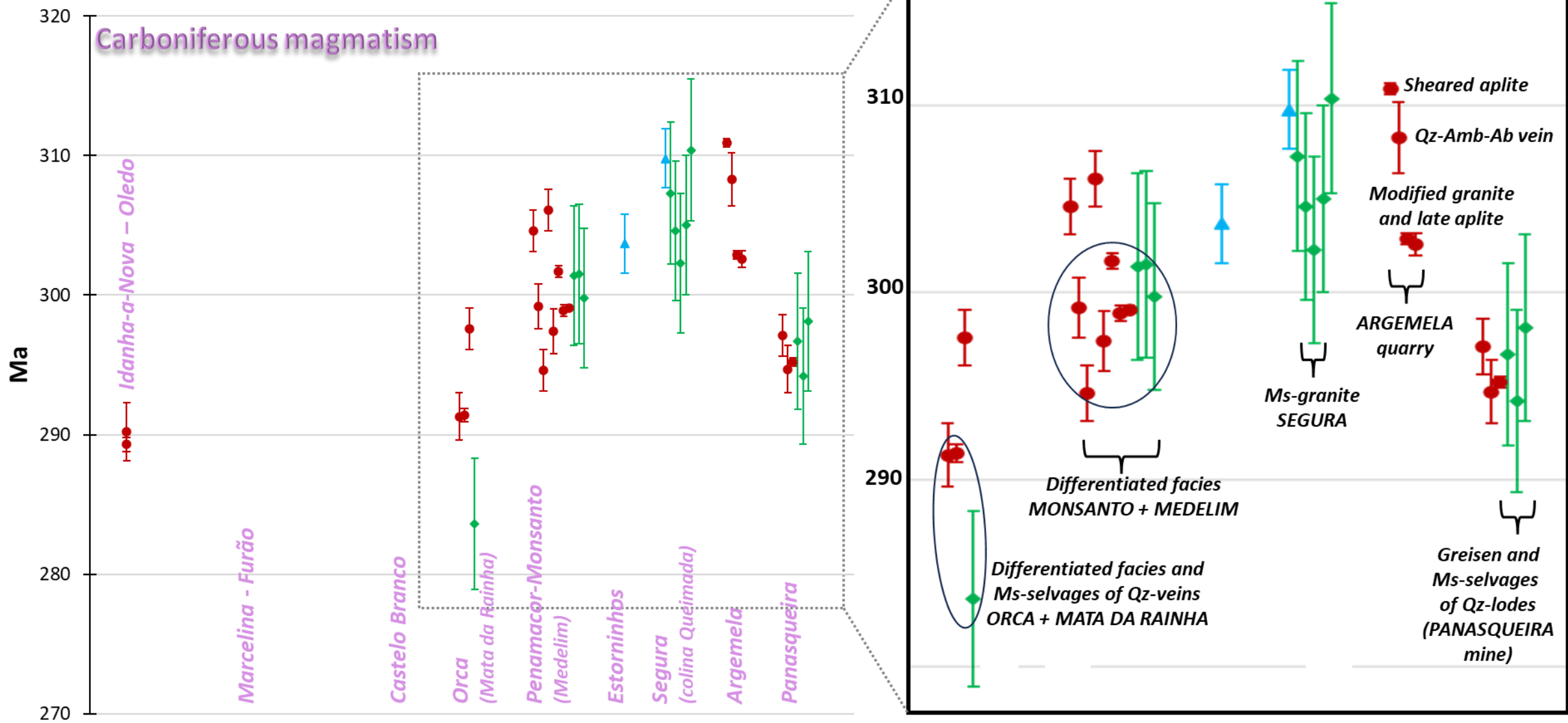
Ar/Ar muscovite plateau age ( $308.3 \pm 1.9 \text{ Ma}$ ) for muscovite extracted from one amblygonite-bearing quartz vein at Cabeço de Argemela (Qz-f-ARG#2)



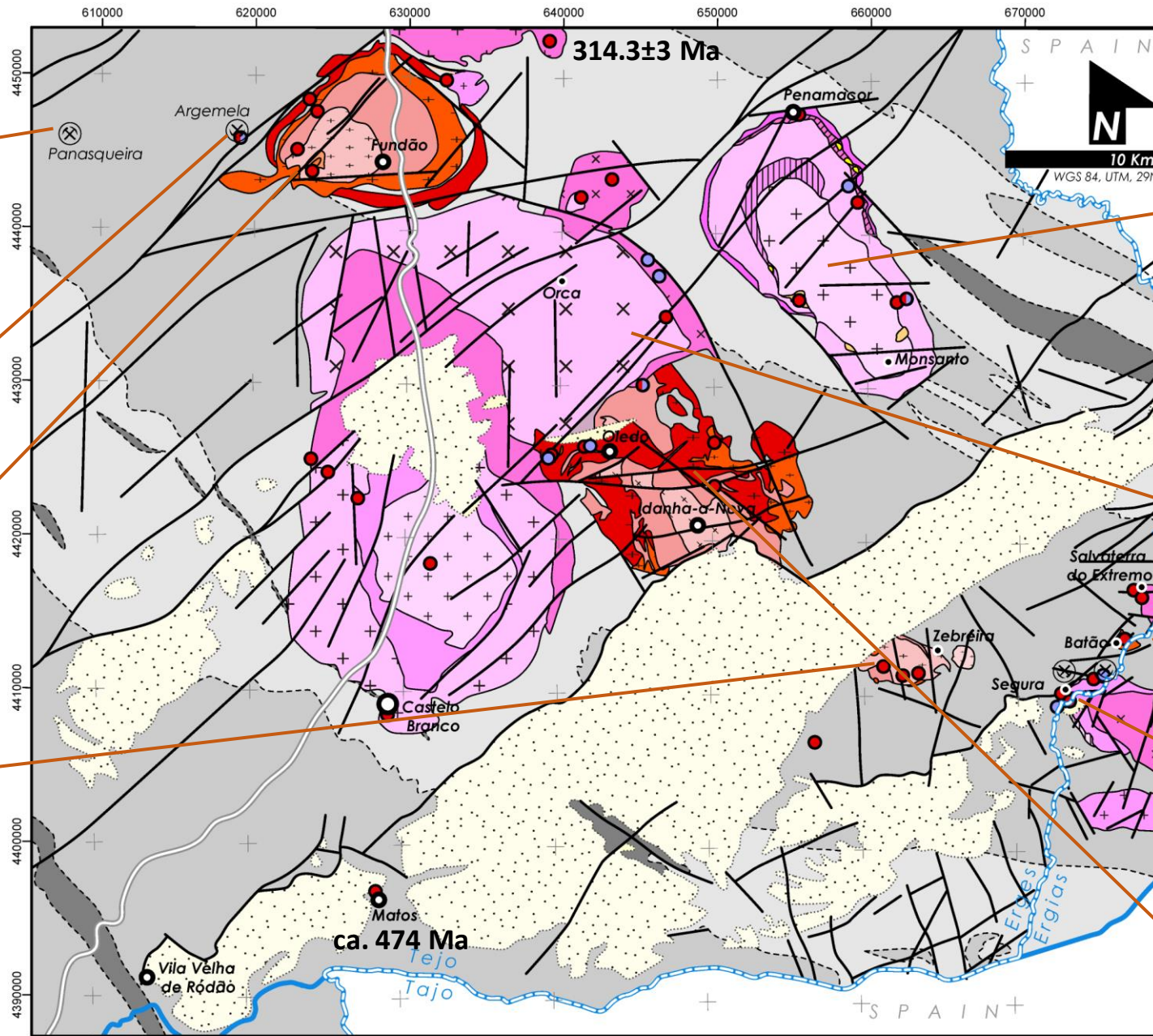
Ar/Ar muscovite (plateau) ages for the pegmatite body (G-MONS#1A,  $306.1 \pm 1.5 \text{ Ma}$ ;



# U-Pb LA-ICP-MS zircon, **Ar-Ar** mica (*ms*), **K-Ar** mica (*ms*)







**PANASQUEIRA**  
 $\gamma$  "two-mica":  
**304±5.8 Ma; 297±1.5 Ma**  
 $\gamma$  "ms": **294.7±1.7 Ma**  
 Greisen: **295.5 – 294.8Ma**

**ARGEMELA**  
 $\gamma$  : **302.6±0.6 Ma**  
 $\gamma$  f: **302.9±0.3 Ma**  
 $\gamma$  f\*: **312 – 310 Ma**

**FUNDÃO**  
 $\gamma^*$ : **499±2 Ma**  
 $\gamma^{**}$ : **498±5 Ma**

**ZEBREIRA**  
 $\gamma$ : **488±3.4 Ma**  
 $\gamma^*$ (dif):  
**468±5Ma; 460.9±2.9 Ma**

**FURÃO/MARCELINA**  
 Prf: **306±4.6 Ma**  
 Prf: **299.9±7.2 Ma**

**PENAMACOR-MONSANTO**  
 $\gamma^*$ : **311.2±2.8 Ma**  
 $\gamma^{**}$ : **308±3 Ma; 304.6±1.9 Ma**  
 $\gamma^{***}$ : **294.6±1.5 Ma**  
 Peg: **306±1.5 Ma**  
 Qz-T: **297.4±1.2 Ma**

**Medelim**  
 $\gamma$  (b): **298.9±1.2 Ma**  
 $\gamma$  (bf): **299.08±0.3 Ma**  
 $\gamma$  f: **302 – 301 Ma**

**ORCA**  
 $\gamma$  : **310±3 Ma**  
 $\gamma$  (b): **304±3 Ma; 291±1.5 Ma**  
 $\gamma$  f: **305 – 295 Ma**  
 MRainha: **297.6±1.5 Ma**

**BATÃO (ton): 508±7 Ma**

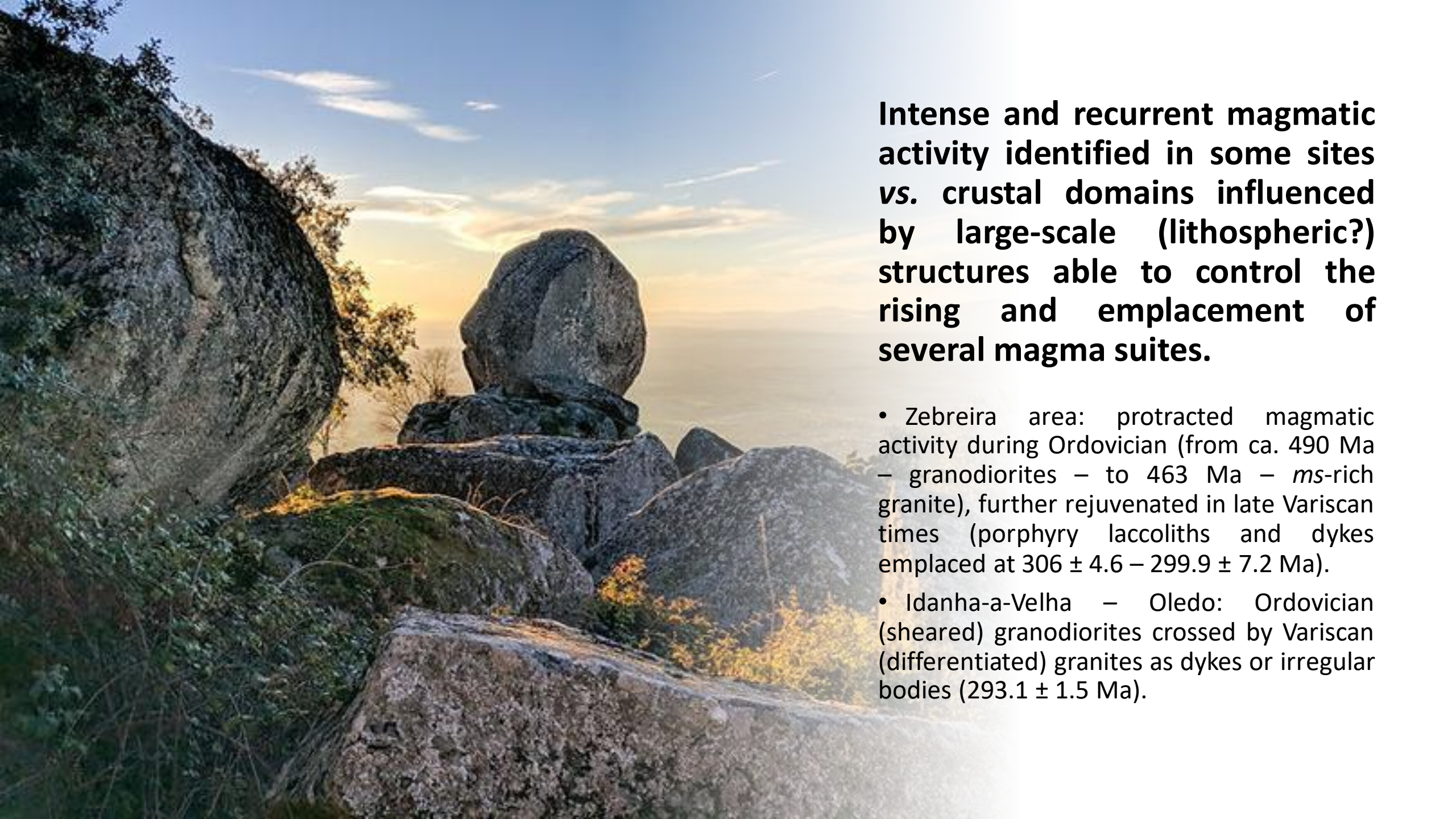
**SEGURA**  
 $\gamma$  : **307.6±2.5 Ma**  
 $\gamma$  ms: **305±2 Ma**  
 Peg: **303±5 Ma**

**IDANHA-A-NOVA-OLEDO**  
 $\gamma$  : **489±4.5 Ma**  
 $\gamma$  f: **293.1±1.5 Ma**  
 $\gamma^*$ (t): **290±2.5 Ma**

**U-Pb in zircon (SHRIMP)**  
**Ar-Ar in muscovite**

⊗ Active mine    ⊗ Abandoned mine    Granite sample location: ● Plutons    ● Dykes





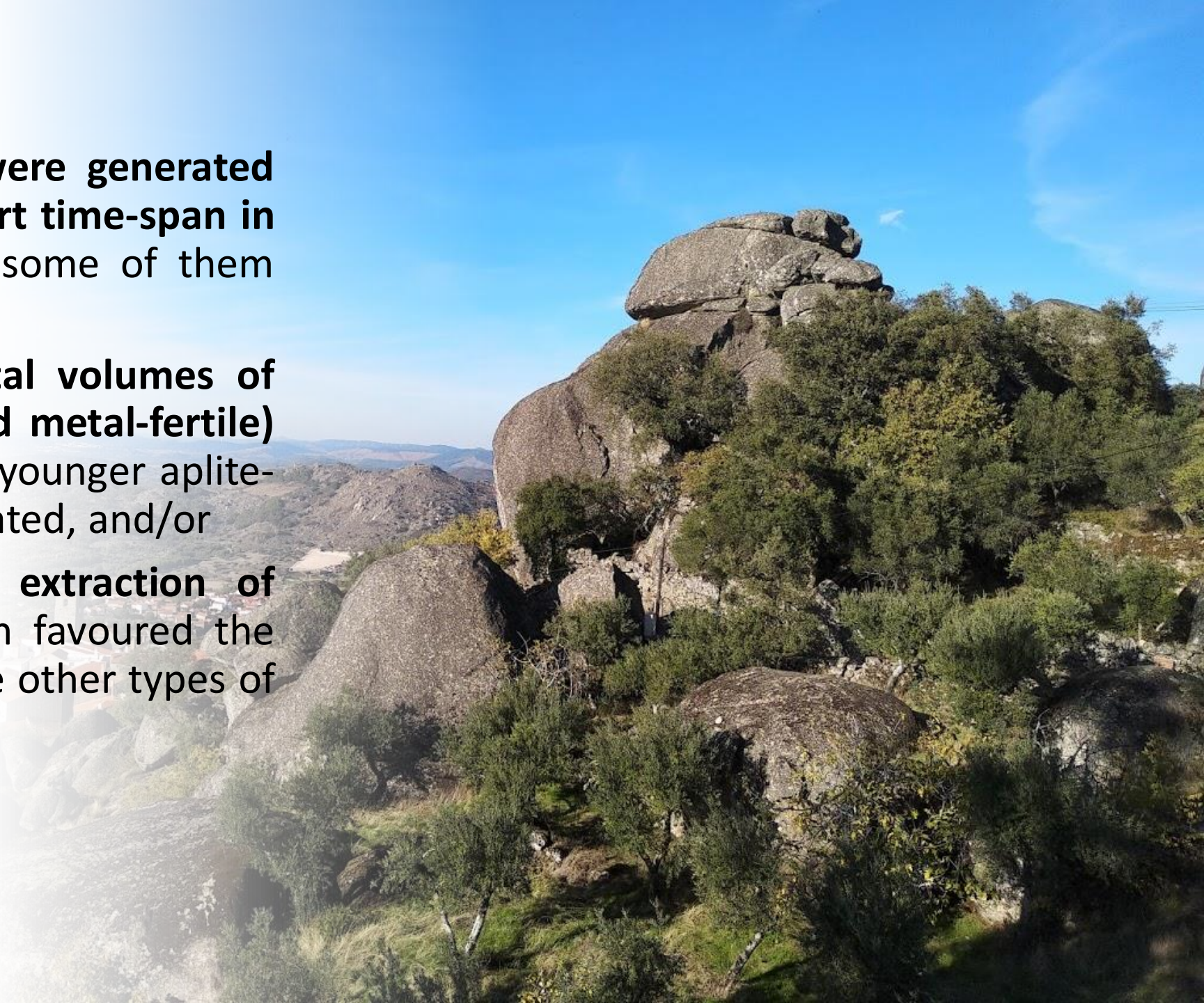
**Intense and recurrent magmatic activity identified in some sites vs. crustal domains influenced by large-scale (lithospheric?) structures able to control the rising and emplacement of several magma suites.**

- Zebreira area: protracted magmatic activity during Ordovician (from ca. 490 Ma – granodiorites – to 463 Ma – *ms*-rich granite), further rejuvenated in late Variscan times (porphyry laccoliths and dykes emplaced at  $306 \pm 4.6$  –  $299.9 \pm 7.2$  Ma).
- Idanha-a-Velha – Oledo: Ordovician (sheared) granodiorites crossed by Variscan (differentiated) granites as dykes or irregular bodies ( $293.1 \pm 1.5$  Ma).



**Successive granite melts were generated and emplaced during a short time-span in the Carboniferous period, some of them followed by the:**

- **Production of incremental volumes of highly differentiated (and metal-fertile) batches** to which several younger aplite-pegmatite swarms are related, and/or
- **Separation and further extraction of supercritical fluids**, which favoured the late development of some other types of mineralised bodies.







**Discrimination of various mineralising events throughout the 310-290 Ma time window might be possible,** separating W-dominant systems from those characterised by the prevalence of Sn and/or Li.

The validation of this interpretation, impacting on the design of future mineral exploration surveys across the strip under study and other similar Variscan orogenic segments, requires additional support.





<https://mostmeg.rd.ciencias.ulisboa.pt/>

Thank you for your attention!

*Segura-Cabeza de Araya pluton (Erges canyon, Segura)*